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### Key Points:

- Reducing particulate nitrate pollution requires understanding its local sensitivities to NH<sub>3</sub>, NO<sub>x</sub>, and volatile organic compound emissions
- Satellite observation of the NH<sub>3</sub>/NO<sub>2</sub> column ratio is an effective indicator for diagnosing these sensitivities
- IASI NH<sub>3</sub> and OMI NO<sub>2</sub> observations reveal varying regimes of nitrate sensitivity across wintertime East Asia

### Supporting Information:

Supporting Information may be found in the online version of this article.

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## Diagnosing the Sensitivity of Particulate Nitrate to Precursor Emissions Using Satellite Observations of Ammonia and Nitrogen Dioxide

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**Abstract** Particulate nitrate is a major component of fine particulate matter (PM<sub>2.5</sub>). Its formation may be varyingly sensitive to emissions of ammonia (NH<sub>3</sub>), nitrogen oxides (NO<sub>x</sub> ≡ NO + NO<sub>2</sub>), and volatile organic compounds (VOCs), depending on local conditions. Diagnosing these sensitivities is critical for successful air quality management. Here, we show that satellite measurements of tropospheric NH<sub>3</sub> and NO<sub>2</sub> columns can be used as a quick indicator of the dominant sensitivity regime through the NH<sub>3</sub>/NO<sub>2</sub> column ratio together with the NO<sub>2</sub> column. We demonstrate the effectiveness of this indicator with the GEOS-Chem chemical transport model and define thresholds to separate the different sensitivity regimes. Applying the method to wintertime IASI and OMI observations in East Asia reveals that surface nitrate is dominantly VOC-sensitive in the southern North China Plain (NCP), NO<sub>x</sub>-sensitive in most of the East China Plain, and NH<sub>3</sub>-sensitive in the northern NCP, southern China, and Korea.

**Plain Language Summary** We present a novel application of satellite remote sensing to better understand the causes of particulate nitrate pollution. Particulate nitrate is a major air pollutant throughout the urbanized world. It is produced by atmospheric oxidation of emitted nitrogen oxides (NO<sub>x</sub>) but may be more sensitive to emissions of ammonia (NH<sub>3</sub>) or volatile organic compounds (VOCs). Understanding which of NH<sub>3</sub>, NO<sub>x</sub>, or VOC emissions is most important in driving nitrate formation is critical for air quality management. We show that satellite measurements of the NH<sub>3</sub>/NO<sub>2</sub> column ratio along with NO<sub>2</sub> columns is an effective indicator to determine the dominant sensitivity regime (NH<sub>3</sub>-, NO<sub>x</sub>-, or VOC - sensitive). We develop this approach using an atmospheric chemistry model and apply it to wintertime satellite observations in East Asia. The approach should be applicable to other continents, seasons, and a broader range of satellite instruments, providing valuable insights for particulate nitrate reduction strategies.

## 1. Introduction

Particulate nitrate (pNO<sub>3</sub><sup>-</sup>) is a major component of fine particulate matter (PM<sub>2.5</sub>) throughout the urbanized world and particularly in winter. It drives PM<sub>2.5</sub> pollution events in East Asia (Li et al., 2018; H. Kim et al., 2020; Kim et al., 2022; Tian et al., 2019; Q. Xu et al., 2019), North America (Franchin et al., 2018; Womack et al., 2019), and Europe (Bressi et al., 2021). It is becoming relatively more important as other PM<sub>2.5</sub> components have decreased in response to emission controls (Attwood et al., 2014; Zhai et al., 2019), but pNO<sub>3</sub><sup>-</sup> has not (Leung et al., 2020; Li et al., 2019; Zhai et al., 2023; Zhou et al., 2022). In eastern China, wintertime pNO<sub>3</sub><sup>-</sup> concentrations have been flat over the past decade despite a 30% decrease in NO<sub>x</sub> emissions (Chuang et al., 2021; Fu et al., 2020; Zhai et al., 2021). pNO<sub>3</sub><sup>-</sup> has become a key target for further improving PM<sub>2.5</sub> air quality.

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$\text{pNO}_3^-$  is produced by the oxidation of nitrogen oxide radicals ( $\text{NO}_x \equiv \text{NO} + \text{NO}_2$ ) to nitric acid ( $\text{HNO}_3$ ).  $\text{HNO}_3$  partitions into the aerosol as  $\text{pNO}_3^-$  depending on aerosol pH, water content, and temperature (Guo et al., 2018; Nenes et al., 2020). The presence of alkalinity, mostly from ammonia ( $\text{NH}_3$ ), raises aerosol pH to favor  $\text{pNO}_3^-$  formation. The resulting  $\text{pNO}_3^-$  is mainly in the fine  $\text{PM}_{2.5}$  aerosol mode.  $\text{NO}_x$  in urban areas mainly comes from fuel combustion.  $\text{NH}_3$  originates from agricultural activities including fertilizer use and livestock manure, but vehicle emissions could also be important in urban areas (Farren et al., 2020; Y. Wang et al., 2023). Oxidation of  $\text{NO}_x$  to  $\text{HNO}_3$  is by the hydroxyl radical ( $\text{OH}$ ) during the daytime and by ozone ( $\text{O}_3$ ) at night, both of which depend on the levels of  $\text{NO}_x$  and volatile organic compounds (VOCs). VOCs originate from combustion, industrial and domestic chemical products, vegetation, and open fires (Shen et al., 2019).

$\text{pNO}_3^-$  concentrations are generally highest in winter when low temperatures favor partitioning into the aerosol. Formation of  $\text{pNO}_3^-$  may then be dominantly sensitive to the gas in shortest supply, either  $\text{NH}_3$  or  $\text{HNO}_3$  (Nenes et al., 2020), while the  $\text{NO}_x$  to  $\text{HNO}_3$  conversion is limited by either the abundance of  $\text{NO}_x$  or VOCs (Kleinman, 1994; Womack et al., 2019). Other factors can further complicate these relationships of  $\text{pNO}_3^-$  to emitted precursors, including competing deposition between  $\text{HNO}_3$  and  $\text{pNO}_3^-$  (Zhai et al., 2021), other sources of aerosol alkalinity (Guo et al., 2018), and  $\text{NO}_x$  oxidation to organic nitrates (Romer Present et al., 2020). Coarse  $\text{pNO}_3^-$  can also form from uptake of  $\text{HNO}_3$  by alkaline soil dust and sea salt in dusty and coastal areas (Zhai et al., 2023). The sensitivity of  $\text{pNO}_3^-$  concentrations to  $\text{NH}_3$ ,  $\text{NO}_x$ , and VOC emissions is thus nonlinear and complex, requiring different control strategies under different conditions.

Two approaches have been used to determine the sensitivity of  $\text{pNO}_3^-$  to emissions. Field studies measure aerosols and gases, allowing for the calculation of diagnostic indicators (Petetin et al., 2016; Z. Xu et al., 2019), or providing input to thermodynamic models for sensitivity tests (Franchin et al., 2018; Guo et al., 2018). They require substantial experimental resources, and the results are only locally applicable. Chemical transport models diagnose the sensitivity of  $\text{pNO}_3^-$  to emissions through simulations with perturbed emissions (Fu et al., 2020; Li et al., 2021; Zhai et al., 2021). They require substantial computational resources, and emission errors in the model may lead to misdiagnosis.

Here we present a new satellite-based method to diagnose locally the sensitivities of fine  $\text{pNO}_3^-$  formation to  $\text{NH}_3$ ,  $\text{NO}_x$ , and VOC emissions. We use for this purpose satellite measurements of the tropospheric column concentrations of  $\text{NH}_3$  ( $\Omega_{\text{NH}_3}$ ) and  $\text{NO}_2$  ( $\Omega_{\text{NO}_2}$ ) and diagnose the sensitivity from the  $\Omega_{\text{NH}_3}/\Omega_{\text{NO}_2}$  ratio. Our approach parallels the common use of the space-based formaldehyde  $\text{HCHO}/\text{NO}_2$  column ratio as an indicator for whether  $\text{O}_3$  formation is  $\text{NO}_x$ - or VOC-limited (Duncan et al., 2010; Jin et al., 2020; Martin et al., 2004). It offers a quick diagnostic tool for air quality management in their design of  $\text{pNO}_3^-$  control strategies.  $\text{NH}_3$  measurements from space have been available from the IASI instrument since 2007 (Clarisse et al., 2009), and from the CrIS instrument since 2012 (Shephard and Cady-Pereira, 2015).  $\text{NO}_2$  measurements from space began with the GOME instrument in 1995 (Martin et al., 2002) and have continued with the OMI instrument since 2005 (Lamsal et al., 2021), the TROPOMI instrument since 2017 (van Geffen et al., 2020), and the GEMS geostationary instrument since 2020 (J. Kim et al., 2020). We demonstrate the method for East Asia in winter, using observations from OMI and IASI.

## 2. Theoretical Basis: $\Omega_{\text{NH}_3}$ and $\Omega_{\text{NO}_2}$ as Indicators of Nitrate Formation Regime

The main pathway for fine  $\text{pNO}_3^-$  formation is the joint condensation of  $\text{NH}_3$  and  $\text{HNO}_3$ , governed by a thermodynamic equilibrium constant  $K$  dependent on temperature and relative humidity (RH) (Stelson and Seinfeld, 1982):

$$K = p_{\text{NH}_3} \times p_{\text{HNO}_3}, \quad (1)$$

where  $p$  is partial pressure. At low wintertime temperatures and/or high RH, the low value of  $K$  leads to titration where  $\text{pNO}_3^-$  formation is mainly sensitive to the gas in shortest supply, either  $\text{NH}_3$  or  $\text{HNO}_3$ . At warmer temperatures,  $\text{NH}_3$  and  $\text{HNO}_3$  may coexist in the gas phase but the dominant sensitivity is still to the gas in shortest supply (Nenes et al., 2020). Scavenging of  $\text{NH}_3$  by acid sulfate may totally suppress  $\text{pNO}_3^-$  formation when sulfate is in excess of  $\text{NH}_3$  (Ansari and Pandis, 1998). Dust and sea salt particles can also drive  $\text{HNO}_3$  into the aerosol through added alkalinity or chloride displacement (Alexander et al., 2005; Fairlie et al., 2010), though this tends to be in coarser particles than  $\text{PM}_{2.5}$ . Henceforth we will refer to  $\text{pNO}_3^-$  as the fine  $\text{PM}_{2.5}$  component of nitrate.

Several frameworks exist for determining the dominant sensitivities in thermodynamic formation of  $\text{pNO}_3^-$ . Nenes et al. (2020) pointed out that aerosol pH is the key variable affecting the dominant  $\text{pNO}_3^-$  sensitivity to  $\text{NH}_3$  or  $\text{HNO}_3$ , and the pH thresholds for distinguishing between the regimes depend on temperature and aerosol

liquid water content. In a more simplified framework, the molar ratio  $R$  of free ammonia after sulfate neutralization ( $\text{NH}_3 + \text{NH}_4^+ - 2 \times \text{SO}_4^{2-}$ ) to total nitrate ( $\text{NO}_3^{\text{T}} \equiv \text{HNO}_3 + \text{pNO}_3^-$ ) is measured in field campaigns to diagnose the sensitivities (Petetin et al., 2016; Z. Xu et al., 2019). The effect of pH is implicitly considered within this simplified indicator through the role of  $\text{NH}_3$  (Guo et al., 2017, 2018). Generally,  $R > 1$  indicates dominant sensitivity to  $\text{HNO}_3$ , while  $R < 1$  indicates dominant sensitivity to  $\text{NH}_3$ . The gas-phase  $\text{NH}_3/\text{HNO}_3$  ratio can also serve as an indicator but its threshold for transition between regimes may depart from unity when  $\text{NO}_3^{\text{T}}$  is heavily partitioned into the aerosol and the resulting  $\text{HNO}_3$  concentration is very low. A dominant sensitivity to  $\text{HNO}_3$  would be expected to translate into a dominant sensitivity to  $\text{NO}_x$  emissions, but the conversion of  $\text{NO}_x$  to  $\text{HNO}_3$  may in fact be limited by the supply of VOCs under VOC-limited conditions for oxidant ( $\text{OH}$  and  $\text{O}_3$ ) formation. Womack et al. (2019) point out that this may cause  $\text{pNO}_3^-$  formation to be most sensitive to VOC emissions under strongly VOC-limited conditions as frequently occur in urban environments in winter.

Satellites measure tropospheric columns of  $\text{NH}_3$  ( $\Omega_{\text{NH}_3}$ ) and  $\text{NO}_2$  ( $\Omega_{\text{NO}_2}$ ). It follows from the above discussion that the measured  $\Omega_{\text{NH}_3}/\Omega_{\text{NO}_2}$  ratio should give an indicator of the sensitivity of  $\text{pNO}_3^-$  formation to precursor emissions, in a manner useful to air quality management. Application of this indicator may be complicated by the vertical gradients of  $\text{NH}_3$  and  $\text{NO}_2$  concentrations, by the presence of sulfate, and by the limiting regime for oxidation of  $\text{NO}_x$  to  $\text{HNO}_3$ . A model analysis can evaluate these complications, and this is discussed in the next section. Satellite observations of  $\text{HCHO}$  columns ( $\Omega_{\text{HCHO}}$ ) could in principle distinguish between  $\text{NO}_x$ - and VOC-limited oxidant regimes through consideration of the  $\Omega_{\text{HCHO}}/\Omega_{\text{NO}_2}$  ratio, but in practice wintertime  $\Omega_{\text{HCHO}}$  concentrations are near or below the detection limit (Zhu et al., 2014, 2017). Very high  $\Omega_{\text{NO}_2}$  values can be used instead as an indicator of VOC-limited conditions (Sillman, 1995).

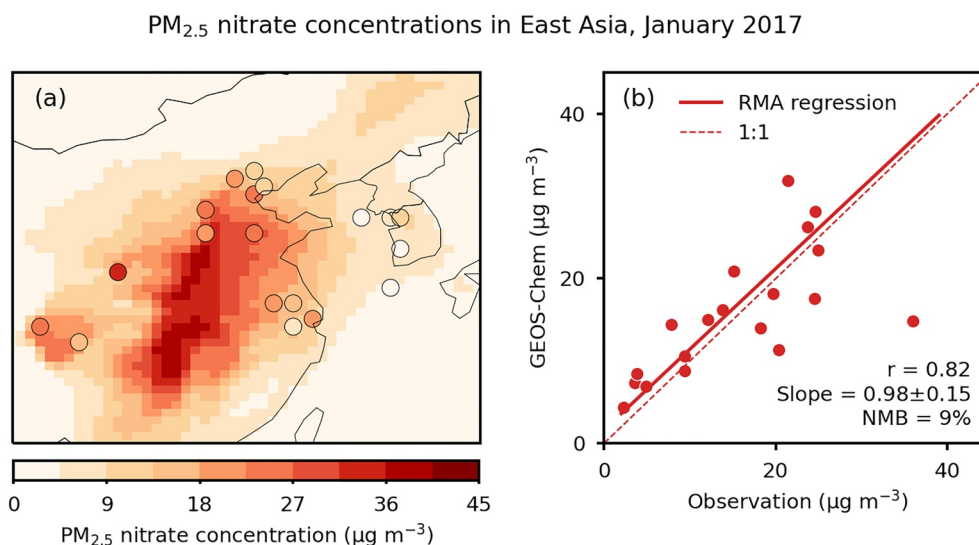
### 3. Evaluation in the GEOS-Chem Model Environment

To analyze the value of the  $\Omega_{\text{NH}_3}/\Omega_{\text{NO}_2}$  ratio as an indicator for the sensitivity of  $\text{pNO}_3^-$  formation to emissions, we conduct sensitivity simulations with the GEOS-Chem global atmospheric chemistry model. We use GEOS-Chem version 13.4.1 (DOI: <https://zenodo.org/record/6564702>) with options and modifications described below. The simulations are driven by MERRA-2 meteorology and are conducted at a nested resolution of  $0.5^\circ \times 0.625^\circ$  for East Asia ( $90^\circ\text{--}145^\circ\text{E}$ ,  $10^\circ\text{--}55^\circ\text{N}$ ) over the 1–31 January 2017 period, with boundary conditions updated every 3 hr from a  $4^\circ \times 5^\circ$  global simulation. The simulation is spun up for 6 months for initialization.

GEOS-Chem includes detailed oxidant-aerosol chemistry (Wang et al., 2021). Thermodynamic  $\text{pNO}_3^-$  formation from  $\text{NH}_3\text{--HNO}_3\text{--H}_2\text{SO}_4\text{--HCl}$  mixtures is calculated by ISORROPIA II (Fountoukis and Nenes, 2007) and defines in the model the  $\text{PM}_{2.5}$  component of  $\text{pNO}_3^-$ . The model also includes uptake of  $\text{HNO}_3$  by coarse sea salt aerosol (Wang et al., 2021) but this does not contribute to  $\text{PM}_{2.5}$  and is not considered here in  $\text{pNO}_3^-$  accounting. Uptake of  $\text{HNO}_3$  by dust is included in GEOS-Chem as an option (Fairlie et al., 2010; Zhai et al., 2023) but is not used in our simulation. We use the wet deposition scheme of Luo et al. (2020), which is an option in GEOS-Chem and has proven to be important for successful simulation of  $\text{pNO}_3^-$  (Luo et al., 2019, 2020; Zhai et al., 2021). We also add to our simulation the photolysis of aerosol nitrate, which improves the simulation of tropospheric  $\text{NO}_2$  column observations in GEOS-Chem though the effect is small in winter (Dang, Jacob, Shah, et al., 2023; Shah et al., 2023). Global anthropogenic emissions are from the Community Emissions Data System (CEDS) (McDuffie et al., 2020) superseded by the MEIC inventory for China (Zheng et al., 2018) and the KORUSv5 inventory for South Korea. Other emissions settings are as described in Dang, Jacob, Shah, et al. (2023).

Figure 1 compares simulated  $\text{pNO}_3^-$  concentrations from our simulation with measurements from observational networks and field studies in China and Korea in winter 2016–2017. Table S1 in Supporting Information S1 gives site details. Most observations are centered on January 2017, but some are for December 2016, and some are for the whole winter (DJF). All are compared to GEOS-Chem in January 2017. GEOS-Chem simulates the ensemble observations with a correlation coefficient  $r = 0.82$ , a reduced-major-axis (RMA) regression slope of  $0.98 \pm 0.15$ , and a normalized mean bias (NMB) of 9%. There is one site in Xi'an where observed  $\text{pNO}_3^-$  is anomalously high (averaging  $36 \mu\text{g m}^{-3}$ ) and this is not captured by the model. This site is excluded from the statistics above. Additionally, GEOS-Chem has been found to reproduce daily  $\text{pNO}_3^-$  observations well at a Beijing site during the winters of 2014–2019 (Zhai et al., 2021). Overall, the successful simulation of  $\text{pNO}_3^-$  variability provides support for using the model to study the sensitivity of  $\text{pNO}_3^-$  to precursor emissions in East Asia.

We diagnose the local  $\text{pNO}_3^-$  sensitivity to  $\text{NH}_3$ ,  $\text{NO}_x$ , and VOC emissions in the model by conducting sensitivity simulations with individual emissions reduced by 20%. The reduction is applied to all sources (anthropogenic



**Figure 1.** Surface  $\text{PM}_{2.5}$  nitrate concentrations in China and Korea. Mean GEOS-Chem model concentrations for January 2017 are compared to mean observations at a number of sites (Table S1 in Supporting Information S1) over December–February 2017. Panel (a) shows the spatial distribution, with observations as circles and GEOS-Chem as solid contours. Panel (b) shows the correlation between model and observations at individual sites including correlation coefficient ( $r$ ), normalized mean bias (NMB), reduced-major-axis (RMA) regression line and slope ( $\pm 95\%$  confidence interval), and 1:1 dashed line. The statistics excludes the Xi’an site where observed  $\text{pNO}_3^-$  is anomalously high. Site details are in Table S1 in Supporting Information S1.

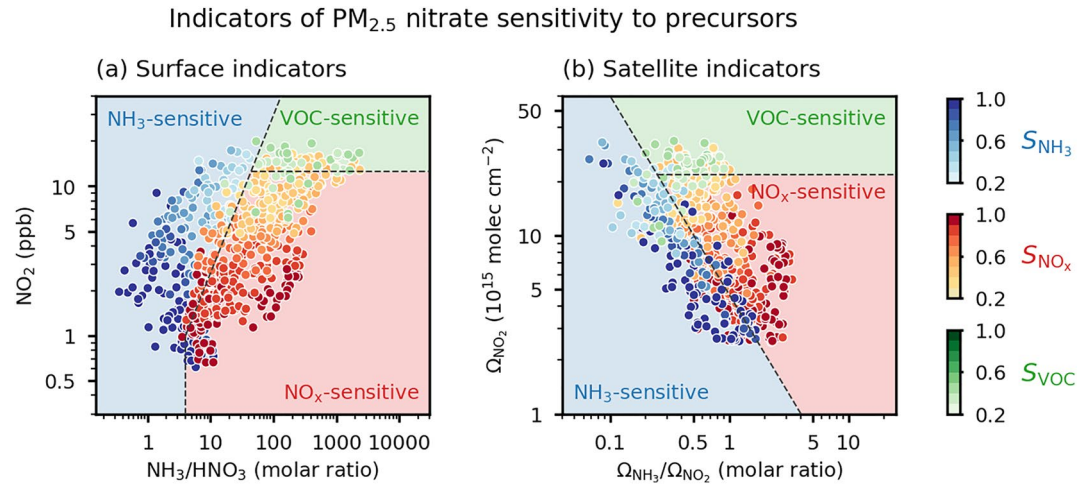
and natural) but the sources in winter are mainly anthropogenic. The local model sensitivity  $S_i$  of  $\text{pNO}_3^-$  to the emission  $E_i$  of species  $i$  for individual  $0.5^\circ \times 0.625^\circ$  grid cells is calculated from the relative model differences ( $\Delta$ ) between the sensitivity and base simulations as:

$$S_i = \frac{\Delta \log[\text{pNO}_3^-]}{\Delta \log E_i}, \quad (2)$$

where  $i$  refers to  $\text{NH}_3$ ,  $\text{NO}_x$ , or VOC, and  $[\text{pNO}_3^-]$  refers to monthly mean concentrations in surface air. A sensitivity  $S_i = 1$  indicates that a 20% reduction in emissions of precursor  $i$  results in a corresponding 20% decrease in surface  $\text{pNO}_3^-$  concentrations. By comparing  $S_{\text{NH}_3}$ ,  $S_{\text{NO}_x}$ , and  $S_{\text{VOC}}$ , we determine whether  $\text{pNO}_3^-$  in a model grid cell is most  $\text{NH}_3$ -,  $\text{NO}_x$ -, or VOC-sensitive.

Figure 2 shows the model relationship between the dominant  $\text{pNO}_3^-$  sensitivity and the observable surface and satellite indicators discussed in Section 2. Individual circles show the dominant sensitivities  $S_i$  for monthly mean surface  $\text{pNO}_3^-$  concentrations in individual grid cells. We use the  $\text{NH}_3/\text{HNO}_3$  gas-phase molar ratio as surface indicator instead of  $R$  because it is better connected to the  $\Omega_{\text{NH}_3}/\Omega_{\text{NO}_2}$  satellite indicator. We use  $\text{NO}_2$  concentration (surface or column) as an indicator of VOC-limited conditions for  $\text{NO}_x$  oxidation because  $\Omega_{\text{HCHO}}$  is generally not observable from space in winter. Surface indicators are 24-hr averages, while columns are sampled at 9–10 local time (LT) for  $\text{NH}_3$  to emulate IASI and at 13–14 LT for  $\text{NO}_2$  to emulate OMI. Averaging kernels are applied to the model  $\text{NO}_2$  vertical profiles following Cooper et al. (2020) to emulate tropospheric  $\text{NO}_2$  columns from version 4 of the NASA OMI  $\text{NO}_2$  level 2 product (OMNO2) (Lamsal et al., 2021). We restrict our attention to grid cells with  $\Omega_{\text{NO}_2} > 2.5 \times 10^{15}$  molec  $\text{cm}^{-2}$  to remove remote regions (as shown by the satellite observations in Figure 3b) where diagnosing sensitivity to local emissions would be inappropriate.

Results in Figure 2 show that the indicators are successful at diagnosing the dominant  $\text{pNO}_3^-$  sensitivities to precursor emissions. Approximately 90% of the grid cells show a dominant sensitivity  $S_i$  that is distinctly greater than the other two sensitivities ( $S_i/S_j > 1.1$ ). Black dashed lines delineate the transitions between sensitivity regimes. The slanted lines are derived from reduced-major-axis (RMA) linear regressions for grid cells with sensitivity ratios  $0.95 < S_i/S_j < 1.05$ . Sensitivities  $S_{\text{NH}_3}$  and  $S_{\text{NO}_x}$  can approach unity within the corresponding regimes.  $S_{\text{VOC}}$  can reach 0.5 in the VOC-sensitive regime.



**Figure 2.** Regimes for the sensitivity of surface  $\text{pNO}_3^-$  concentrations to  $\text{NH}_3$ ,  $\text{NO}_x$ , and VOC emissions. Results show the dominant sensitivities  $S_i = \Delta \log[\text{pNO}_3^-] / \Delta \log E_i$  for monthly mean concentrations in January 2017 in individual  $0.5^\circ \times 0.625^\circ$  GEOS-Chem model grid cells in East Asia (domain of Figure 1(a)). A sensitivity  $S_i = 1$  indicates proportional response of the  $\text{pNO}_3^-$  concentration to change in the precursor emission  $E_i$ . The dominant sensitivities are plotted in a state space of indicators of the sensitivity regime as observable from surface or satellite measurements. Surface indicators (panel (a)) are the gas-phase  $\text{NH}_3/\text{HNO}_3$  molar ratio and the  $\text{NO}_2$  concentration. Satellite indicators (panel (b)) are the  $\Omega_{\text{NH}_3}/\Omega_{\text{NO}_2}$  column ratio and the  $\Omega_{\text{NO}_2}$  column.  $\Omega_{\text{NH}_3}$  is sampled at 9–10 local time (LT) to emulate the IASI instrument, and  $\Omega_{\text{NO}_2}$  is sampled at 13–14 LT to emulate the OMI instrument. Dashed lines separate the different regimes as diagnosed by  $S_i$ . The slanted lines are derived from reduced-major-axis (RMA) linear regression for grid cells with sensitivity ratios  $0.95 < S_i/S_j < 1.05$ . The corresponding equations are given in the text. Ocean and remote grid cells with  $\Omega_{\text{NO}_2} < 2.5 \times 10^{15} \text{ molec cm}^{-2}$  (see Figure 3b) are excluded from the plot.

Examining first the surface indicators, we find that  $\text{NH}_3$ -sensitive conditions are associated with  $\text{NH}_3/\text{HNO}_3 < 4 \text{ mol mol}^{-1}$  at low  $\text{NO}_2$ , with the threshold increasing at higher  $\text{NO}_2$ . The threshold is larger than the value of 1 for the  $R$  ratio in Section 2. This is expected because the gas-phase  $\text{HNO}_3$  concentration can be extremely low in winter, so that competing deposition between gas-phase  $\text{HNO}_3$  and  $\text{pNO}_3^-$  increases sensitivity to  $\text{NH}_3$  even when  $R > 1$  (Zhai et al., 2021). Outside of the  $\text{NH}_3$ -sensitive regime, whether  $\text{NO}_x$  or VOCs is the controlling precursor is well delineated by  $\text{NO}_2$  levels. For  $\text{NO}_2 < 12 \text{ ppb}$  the sensitivity is mostly to  $\text{NO}_x$  emissions ( $\text{NO}_x$ -limited regime) but it decreases as  $\text{NO}_2$  increases and VOCs then become more important. For  $\text{NO}_2 > 12 \text{ ppb}$  the sensitivity is mostly to VOCs (strongly VOC-limited regime).  $\text{NH}_3$  sensitivity can also be dominant under these conditions because the conversion efficiency of  $\text{NO}_x$  to  $\text{HNO}_3$  is low. The sensitivity regimes separated by the black dashed lines in Figure 2a are thus diagnosed from the gas-phase  $\text{NH}_3/\text{HNO}_3$  and  $\text{NO}_2$  surface indicator concentrations as

$$\text{NH}_3 - \text{sensitive: } \begin{cases} \frac{[\text{NH}_3]}{[\text{HNO}_3]} < 4 & ([\text{NO}_2] < 1.3 \text{ ppb}) \\ \log \frac{[\text{NH}_3]}{[\text{HNO}_3]} < 0.49 + 1.02 \times \log [\text{NO}_2] & ([\text{NO}_2] > 1.3 \text{ ppb}) \end{cases}, \quad (3a)$$

$$\text{NO}_x - \text{sensitive: } \begin{cases} \frac{[\text{NH}_3]}{[\text{HNO}_3]} > 4 & ([\text{NO}_2] < 1.3 \text{ ppb}) \\ \log \frac{[\text{NH}_3]}{[\text{HNO}_3]} > 0.49 + 1.02 \times \log [\text{NO}_2] & (1.3 \text{ ppb} < [\text{NO}_2] < 12 \text{ ppb}) \end{cases}, \quad (3b)$$

$$\text{VOC} - \text{sensitive: } \log \frac{[\text{NH}_3]}{[\text{HNO}_3]} > 0.49 + 1.02 \times \log [\text{NO}_2] \quad ([\text{NO}_2] > 12 \text{ ppb}). \quad (3c)$$

Figure 2b shows that the satellite indicators are similarly effective for diagnosing sensitivity regimes. For a given  $\Omega_{\text{NH}_3}/\Omega_{\text{NO}_2}$  ratio, higher  $\Omega_{\text{NO}_2}$  levels indicate a lower efficiency in converting  $\text{NO}_2$  to  $\text{HNO}_3$ , so that  $\text{NH}_3$

is more likely to be in excess. This explains why the threshold  $\Omega_{\text{NH}_3}/\Omega_{\text{NO}_2}$  ratio for transition from  $\text{NH}_3$ -sensitive to  $\text{NO}_x$ -sensitive conditions decreases with increasing  $\Omega_{\text{NO}_2}$ , while by contrast the threshold  $\text{NH}_3/\text{HNO}_3$  ratio in surface observations increases with increasing  $\text{NO}_2$ . We also see from Figure 2 that  $\Omega_{\text{NO}_2}$  can serve as a good satellite indicator for the onset of VOC-sensitive conditions. The sensitivity regimes separated by the black dashed lines in Figure 2b are thus diagnosed from the  $\Omega_{\text{NH}_3}$  and  $\Omega_{\text{NO}_2}$  columns as

$$\text{NH}_3 - \text{sensitive: } \log \frac{\Omega_{\text{NH}_3}}{\Omega_{\text{NO}_2}} < 14.09 - 0.90 \times \log \Omega_{\text{NO}_2}, \quad (4a)$$

$$\text{NO}_x - \text{sensitive: } \log \frac{\Omega_{\text{NH}_3}}{\Omega_{\text{NO}_2}} > 14.09 - 0.90 \times \log \Omega_{\text{NO}_2} \quad (\Omega_{\text{NO}_2} < 2 \times 10^{16} \text{ molec cm}^{-2}), \quad (4b)$$

$$\text{VOC} - \text{sensitive: } \log \frac{\Omega_{\text{NH}_3}}{\Omega_{\text{NO}_2}} > 14.09 - 0.90 \times \log \Omega_{\text{NO}_2} \quad (\Omega_{\text{NO}_2} > 2 \times 10^{16} \text{ molec cm}^{-2}). \quad (4c)$$

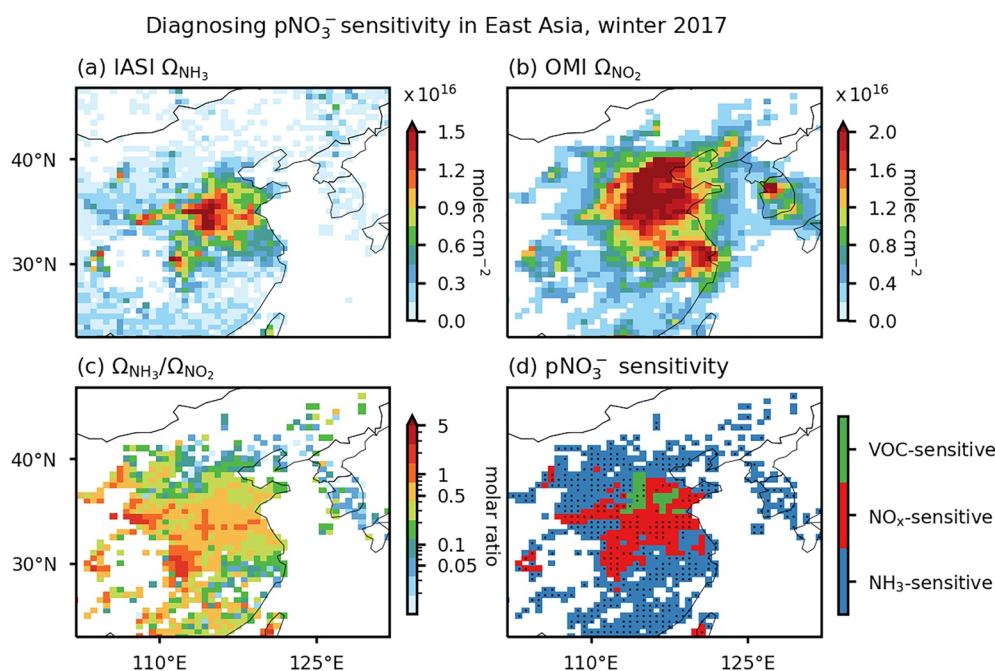
#### 4. Application to Satellite Observations

We now illustrate the application of the method to satellite observations of  $\Omega_{\text{NH}_3}$  from IASI and  $\Omega_{\text{NO}_2}$  from OMI, using Equation 4 to diagnose the sensitivity regimes in the observations. The IASI instrument measures  $\Omega_{\text{NH}_3}$  by observing the infrared radiation emitted by the Earth's surface and the atmosphere. It provides global coverage twice a day, at 9:30 local solar time (LT) and 21:30 LT, with a nadir pixel resolution of  $12 \times 12 \text{ km}^2$  (Van Damme et al., 2014). The OMI instrument measures  $\Omega_{\text{NO}_2}$  by observing solar backscatter, providing daily global coverage at 13:30 LT with a nadir pixel resolution of  $13 \times 24 \text{ km}^2$ . Here, we use version 3 of the reanalyzed level 2 product of  $\text{NH}_3$  columns (ANNI- $\text{NH}_3$ -v3R) (Van Damme et al., 2021) and version 4 of the NASA OMI  $\text{NO}_2$  level 2 product (OMNO2) (Lamsal et al., 2021) during the winter (DJF) of 2017. Both products have been extensively validated including for IASI v3 (Guo et al., 2021; Vohra et al., 2021; Wang et al., 2022; R. Wang et al., 2023) and OMNO2 version 4 (Lamsal et al., 2021). Both data sets have been used effectively in previous studies for hotspot detection (Clarisse et al., 2019; Mebust et al., 2011) and emission tracking (Chen et al., 2021; Cooper et al., 2022; Evangeliou et al., 2021; Luo et al., 2022; Marais et al., 2021; Shah et al., 2020).

We only use morning overpasses (9:30 LT) for  $\Omega_{\text{NH}_3}$  to minimize the time separation with OMI afternoon observations. We filter the IASI  $\Omega_{\text{NH}_3}$  data to remove pixels with cloud fraction  $>0.1$ . For OMI  $\Omega_{\text{NO}_2}$  data, we filter out pixels with cloud fraction  $>0.3$ , surface reflectivity  $>0.3$ , solar zenith angle  $>75^\circ$ , viewing zenith angle  $>65^\circ$ , and those affected by the so-called row anomaly. To reduce noise, both data sets are gridded and averaged to obtain wintertime mean columns at  $0.5^\circ \times 0.625^\circ$  resolution, and grid cells with fewer than 30 successful retrievals for either  $\Omega_{\text{NH}_3}$  or  $\Omega_{\text{NO}_2}$  are excluded. Additional filtering is applied to the gridded wintertime means to remove negative values. Uncertainties in grid-cell averages for both data sets are calculated using the method described by Eskes et al. (2003), with a 0.15 error correlation applied to retrievals falling within the same grid cell (Boersma et al., 2018). The calculated uncertainties range from 14% to 85% (0.1–0.9 quantiles) for IASI  $\Omega_{\text{NH}_3}$  and 9% to 26% for OMI  $\Omega_{\text{NO}_2}$  over the studied region (Figure 3) during the winter of 2017.

Figures 3a and 3b show the IASI  $\Omega_{\text{NH}_3}$  and OMI  $\Omega_{\text{NO}_2}$  during the winter of 2017. IASI observes high  $\text{NH}_3$  in the East China Plain where it originates from livestock waste, fertilizer use, and vehicles (Zhang et al., 2018). OMI observes high  $\text{NO}_2$  in the densely populated East China Plain and the Seoul metropolitan area (SMA) in South Korea. These satellite observations of  $\Omega_{\text{NH}_3}$  and  $\Omega_{\text{NO}_2}$  are roughly consistent with the GEOS-Chem simulations (Figures S1 and S2 in Supporting Information S1) but that is not a requirement for application of our method.

Figure 3d shows the dominant local surface  $\text{pNO}_3^-$  sensitivities to precursor emissions determined from the observed  $\Omega_{\text{NH}_3}/\Omega_{\text{NO}_2}$  ratio (Figure 3c) and  $\Omega_{\text{NO}_2}$  (Figure 3b) by applying Equation 4. We assume that the thresholds obtained from Equation 4 in January can represent the entirety of winter (DJF), considering that the effect of meteorological variability over those 3 months is small compared to the range of conditions within the spatial domain sampled by the model. Robustness tests are conducted for each grid cell by applying 10,000 Monte Carlo samplings for both IASI  $\Omega_{\text{NH}_3}$  and OMI  $\Omega_{\text{NO}_2}$  data, with grid means and uncertainties as inputs to describe the distributions. Grid cells exhibiting a robust diagnosis with a 90% confidence level are marked with black dots.



**Figure 3.** Sensitivity of surface particulate nitrate ( $\text{pNO}_3^-$ ) concentrations in East Asia to precursor emissions as diagnosed from mean satellite observations in winter (DJF) 2016–2017. Panels (a) and (b) show IASI observations of  $\text{NH}_3$  columns ( $\Omega_{\text{NH}_3}$ ) and OMI observations of tropospheric  $\text{NO}_2$  columns ( $\Omega_{\text{NO}_2}$ ), filtered as described in the text. Panel (c) shows the molar  $\Omega_{\text{NH}_3}/\Omega_{\text{NO}_2}$  ratio computed from the seasonal mean columns. Panel (d) presents the dominant sensitivity regimes of  $\text{pNO}_3^-$  diagnosed from the satellite observations using Equation 4. White areas indicate either lack of data or remote areas ( $\Omega_{\text{NO}_2} < 2.5 \times 10^{15}$  molec  $\text{cm}^{-2}$ ). Black dots indicate grid cells with robust diagnoses at a 90% confidence level, determined through 10,000 Monte Carlo samplings for both  $\Omega_{\text{NH}_3}$  and  $\Omega_{\text{NO}_2}$ .

We find varying regimes of  $\text{pNO}_3^-$  sensitivity across China and Korea. VOC-sensitive conditions are observed in the southern North China Plain (NCP), characterized by a  $\Omega_{\text{NH}_3}/\Omega_{\text{NO}_2}$  molar ratio exceeding 0.5 and  $\Omega_{\text{NO}_2}$  exceeding  $2 \times 10^{16}$  molec  $\text{cm}^{-2}$ . In this region,  $\text{pNO}_3^-$  formation is  $\text{NH}_3$ -saturated, and the most effective approach to decrease  $\text{pNO}_3^-$  is to control VOC emissions. In other areas of the East China Plain including Henan and Hubei provinces, and in the Fenwei Plain, the satellite observations indicate  $\text{NO}_x$ -sensitive conditions. In these areas,  $\text{NH}_3$  levels are high and  $\text{NO}_x$  concentrations are not as high as in the southern NCP, so controlling  $\text{NO}_x$  emissions is the most effective way for decreasing  $\text{pNO}_3^-$ .  $\text{NH}_3$ -sensitive conditions are observed in the northern NCP (including Beijing), southern China, and Korea, characterized by relatively low  $\Omega_{\text{NH}_3}/\Omega_{\text{NO}_2}$  ratios.

Previous field studies found that  $\text{pNO}_3^-$  formation at sites in eastern China are more sensitive to total nitrate than to  $\text{NH}_3$  due to  $\text{NH}_3$  being present in excess (Guo et al., 2018; Lin et al., 2020; Song et al., 2019; Zang et al., 2022). However, difference in the lifetimes of  $\text{HNO}_3$  and  $\text{pNO}_3^-$  against deposition can drive a dominant sensitivity to  $\text{NH}_3$  even when  $\text{NH}_3$  is present in excess (Nenes et al., 2021; Zhai et al., 2021), as reflected in our GEOS-Chem simulation where the  $\text{NH}_3$ -sensitive regime extends to  $\text{NH}_3/\text{HNO}_3$  gas-phase ratios in excess of unity (Figure 2a). Our findings are consistent with previous model studies, where wintertime  $\text{pNO}_3^-$  concentrations are found to be most sensitive to  $\text{NH}_3$  and/or VOC controls in the NCP (Fu et al., 2020; Li et al., 2021; Zhai et al., 2021) and to  $\text{NH}_3$  controls in the Yangtze River Delta (Li et al., 2021) and southern China (Lu et al., 2021).

Our demonstration of this satellite-based method for diagnosing the sensitivity of  $\text{pNO}_3^-$  to emissions has focused on wintertime East Asia, where  $\text{pNO}_3^-$  is particularly high. One might expect the same method and similar thresholds to be applicable to other polluted regions and seasons, but this would need to be further investigated with model simulations and evaluated with local field studies.

In summary, we have shown that  $\text{NH}_3$  and  $\text{NO}_2$  measurements from space can be used as a  $\text{NH}_3/\text{NO}_2$  column ratio indicator to diagnose the sensitivity of  $\text{PM}_{2.5}$  nitrate to emissions in support of pollution management. Our method could be applied to other current satellite instruments including TROPOMI for  $\text{NO}_2$  and CrIS for  $\text{NH}_3$ . Future geostationary satellites including Sentinel-4 and IRS for Europe (Gulde et al., 2017) and GeoXO for the

United States (Schmit et al., 2022) will have NO<sub>2</sub> and NH<sub>3</sub> measurements from the same platform. The Nitrosat satellite mission presently under consideration by the European Space Agency will simultaneously observe NH<sub>3</sub> and NO<sub>2</sub> at 500-m resolution, greatly increasing the frequency of clear-sky scenes (Coheur et al., 2021). There is thus considerable potential for application of our method to the next generation of satellite observations. This new satellite-based method would enable us to gain a global perspective on pNO<sub>3</sub><sup>-</sup> sensitivity and monitor regime changes.

## Data Availability Statement

The IASI reanalyzed daily NH<sub>3</sub> data are publicly available from Clarisse et al. (2022). The OMNO2 product, created by the National Aeronautics and Space Administration (NASA), is available at Krotkov et al. (2019). The PM<sub>2.5</sub> nitrate observation data that are collected in this study can be accessed via Dang, Jacob, Zhai, et al. (2023).

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