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JGR Atmospheres

RESEARCH ARTICLE

10.1029/2024JD041784

Key Points:

- Temperature (T-) dependency of biogenic HCHO columns varies substantially across plant functional types (PFTs)
- The GEOS‐Chem model with the MEGAN module implemented primarily interprets the T‐dependency of HCHO columns at the PFT level
- The T-dependency of biogenic volatile organic compound (BVOC) emissions mainly accounts for that of HCHO columns in Broadleaf Evergreen Tropical Trees and Warm C4 Grass regions

Supporting Information:

Supporting Information may be found in the online version of this article.

Correspondence to:

L. Zhu, zhul3@sustech.edu.cn

Citation:

Li, X., Zhu, L., De Smedt, I., Sun, W., Chen, Y., Shu, L., et al. (2024). Global temperature dependency of biogenic HCHO columns observed from space: Interpretation of TROPOMI results using GEOS‐Chem model. *Journal of Geophysical Research: Atmospheres*, *129*, e2024JD041784. [https://doi.org/10.1029/](https://doi.org/10.1029/2024JD041784) [2024JD041784](https://doi.org/10.1029/2024JD041784)

Received 12 JUN 2024 Accepted 21 OCT 2024

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Global Temperature Dependency of Biogenic HCHO Columns Observed From Space: Interpretation of TROPOMI Results Using GEOS‐Chem Model

Xicheng Li¹ (D[,](https://orcid.org/0000-0002-4099-8958) Lei Zhu^{1,2,3} (D, Isabelle De Smedt⁴ (D, Wenfu Sun⁴ (D, Yuyang Chen¹ (D, Lei Shu⁵ (D, **Dakang Wang[6](https://orcid.org/0000-0003-0342-7264) , Song Liu1 [,](https://orcid.org/0000-0002-4063-0205) Dongchuan Pu1 , Juan Li1 , Xiaoxing Zuo¹ , Weitao Fu1 , Yali Li¹ , Peng Zhang¹ , Zhuoxian Yan1 , Tzung‐May Fu[1,2](https://orcid.org/0000-0003-1980-9734),3 [,](https://orcid.org/0000-0002-8556-7326) Huizhong Shen1,2,3 , Chen** $\text{Wang}^{1,2,3}$ $\text{Wang}^{1,2,3}$ $\text{Wang}^{1,2,3}$ **.**, **Jianhuai** $\text{Ye}^{1,2,3}$ **D**, and Xin $\text{Yang}^{1,2,3}$ **D**

1 School of Environmental Science and Engineering, Southern University of Science and Technology, Shenzhen, China, 2 Guangdong Provincial Observation and Research Station for Coastal Atmosphere and Climate of the Greater Bay Area, Shenzhen, China, ³Shenzhen Key Laboratory of Precision Measurement and Early Warning Technology for Urban Environmental Health Risks, School of Environmental Science and Engineering, Southern University of Science and Technology, Shenzhen, China, ⁴Division of Atmospheric Composition, Royal Belgian Institute for Space Aeronomy, Brussels, Belgium, ⁵School of Geographical Sciences, Fujian Normal University, Fuzhou, China, ⁶School of Geography and Remote Sensing, Guangzhou University, Guangzhou, China

Abstract Temperature is the principal driver of global atmospheric formaldehyde (HCHO) and its primary oxidation precursor biogenic volatile organic compounds (BVOCs). We revisit such a temperature (T‐) dependency globally, leveraging TROPOMI HCHO column data. We find substantial variations in the T‐ dependency of biogenic HCHO across plant functional types (PFTs), with the highest over Broadleaf Evergreen Tropical Trees (doubling every $6.0 K \pm 4.1 K$) and lowest over Arctic C3 Grass (doubling every 30.8 K \pm 9.6 K). The GEOS-Chem model interprets HCHO columns' T-dependency at the PFT level ($r = 0.87$), with a 16% discrepancy on average. The discrepancy can be explained by BVOC emissions T-dependency for Broadleaf Evergreen Tropical Trees and Warm C4 Grass and can be attributed to the insensitivity of HCHO columns to BVOC emissions for other PFTs. Our findings underscore a potentially magnified variation of BVOC emissions by GEOS‐Chem and MEGAN therein, particularly in regions experiencing greater temperature variations.

Plain Language Summary We use remote sensing data from an up-to-date monitor to examine the temperature (T‐) dependency of biogenic formaldehyde (HCHO), a proxy of a series of volatile organic gases released by plants, in a global manner. We find that the effect of temperature on HCHO varies significantly between different types of plants, with tropical evergreen trees showing the most sensitivity to temperature and Arctic grasses showing the least. The GEOS-Chem, a state-of-the-art chemical transport model, interprets such temperature sensitivity among plants with nonnegligible discrepancies. The sensitivity of volatile organic gases released by plants to temperature explains the sensitivity of HCHO to temperature for some plants, such as tropical evergreen trees and warm‐season grasses.

1. Introduction

Biogenic volatile organic compounds (BVOCs) account for 90% of the total VOC emissions (Guenther et al., [1995](#page-7-0)), significantly affecting atmospheric chemistry and air quality via the secondary production of tropospheric ozone (Hofzumahaus et al., [2009](#page-7-0); Pacifico et al., [2012\)](#page-8-0) and organic aerosols (Janssen et al., [2013](#page-7-0); Kroll et al., [2006](#page-7-0)). Isoprene is the most abundant BVOC species (Arneth et al., [2008](#page-6-0); Guenther et al., [2012](#page-7-0); Sindelarova et al., [2014\)](#page-8-0) with emissions highly dependent on temperature, as reported in North and South America, North Europe, and Africa (Alves et al., [2018](#page-6-0); Bourtsoukidis et al., [2024;](#page-6-0) Harley et al., [2003](#page-7-0), [2004](#page-7-0); Langford et al., [2017;](#page-7-0) Rhew et al., [2017](#page-8-0); Stoy et al., [2021](#page-9-0)). Previous studies have documented satellite HCHO columns as a proxy of BVOC emissions and demonstrated that regional HCHO columns depend on temperature variations (Duncan et al., [2009](#page-7-0); Kaiser et al., [2018](#page-7-0); Palmer et al., [2006;](#page-8-0) Zhu et al., [2014\)](#page-9-0). Here, we examine the global distribution of such a temperature (T‐) dependency in biogenic‐dominated regions for the first time using high-resolution observations from the TROPOspheric Monitoring Instrument (TROPOMI) (Veefkind et al., [2012\)](#page-9-0). We interpret such T-dependency using MEGAN implemented in GEOS-Chem model.

BVOC emissions are sensitive to temperature accompanied by other environmental factors, including plant function types (PFTs) (Arneth et al., [2011;](#page-6-0) Bonan et al., [2002;](#page-6-0) Bourtsoukidis et al., [2024;](#page-6-0) Misztal et al., [2011](#page-8-0), [2016](#page-8-0)), photosynthesis (Litvak et al., [1996;](#page-8-0) Van Meeningen et al., [2017;](#page-9-0) Wang et al., [2022](#page-9-0)), soil moisture (Naimark et al., [2021;](#page-8-0) Zheng et al., [2017](#page-9-0)), and CO₂ concentrations (Garcia et al., [2019;](#page-7-0) Heald et al., [2009;](#page-7-0) Tai et al., [2013\)](#page-9-0). Although isoprene emission flux has been measured among plant species and environmental con-ditions globally ranging from 0.1 to 93 mg C m⁻² day⁻¹ (Cao et al., [2021\)](#page-6-0), site measurements lack representativeness of the region. Different choices of model parameters as well as the limited understanding of those environmental factors can give results with a factor of four errors for one region (Wu et al., [2020](#page-9-0)). Therefore, accurately quantifying BVOC emissions globally is highly challenging.

As a highly yielded product from BVOC oxidations, HCHO has an atmospheric lifetime of a few hours, making it a proxy of local and global BVOC emissions (Barkley et al., [2013](#page-6-0); Curci et al., [2010;](#page-6-0) Millet et al., [2006,](#page-8-0) [2008](#page-8-0); Palmer et al., [2003,](#page-8-0) [2006](#page-8-0); Shim et al., [2005](#page-8-0); Wolfe et al., [2016;](#page-9-0) Zhu et al., [2014;](#page-9-0) Zhu, & Mickley et al., [2017\)](#page-9-0). Temperature is the most influential driver of satellite HCHO column variations (Abbot et al., [2003](#page-6-0); Duncan et al., [2009;](#page-7-0) Guenther et al., [2006](#page-7-0); Naimark et al., [2021;](#page-8-0) Palmer et al., [2006](#page-8-0); Steiner et al., [2010;](#page-9-0) Zhu et al., [2014](#page-9-0); Zhu, & Mickley et al., [2017](#page-9-0)), likely related to the rate-limiting process of BVOC (especially isoprene) synthase enzymatic reaction in plants (Monson et al., [1992;](#page-8-0) Rasulov et al., [2010;](#page-8-0) Sharkey et al., [2007](#page-8-0)). HCHO columns have been reported to be exponentially dependent on temperature in regions with abundant BVOC emissions (Ryan, & Rhodes et al., [2020](#page-8-0); Ryan, & Silver et al., [2020;](#page-8-0) Zhu et al., [2014\)](#page-9-0). This study explores such dependency globally using a 5‐year monthly TROPOMI HCHO columns over the biogenic‐dominated areas and interprets the dependency with the state-of-the-art MEGAN module v2.1 (Guenther et al., [2012\)](#page-7-0) implemented in the GEOS-Chem model (v12.9.3) (Bey et al., [2001](#page-6-0)).

2. TROPOMI HCHO Columns

HCHO is detectable from space as a vertical column density (VCD), retrieved from the backscattered ultraviolet radiance from 325 to 360 nm (Chance et al., [2000](#page-6-0)). TROPOMI, on board the Copernicus Sentinel‐5 Precursor satellite (S5P) launched in October 2017, is a nadir viewing shortwave spectrometer instrument using passive remote sensing techniques to attain its objectives at 13:30 daily local time. Among the available products, TROPOMI provides retrievals at a finer spatial resolution of 7.0 \times 3.5 km² (upgraded to 5.5 \times 3.5 km² since August 2019) and signal‐to‐noise ratio (De Smedt et al., [2018](#page-6-0), [2021](#page-6-0)), with the algorithm inherited and updated from its predecessors, including the Global Ozone Monitoring Experiment (GOME) (De Smedt et al., [2008\)](#page-6-0), SCanning Imaging Absorption spectroMeter for Atmospheric ChartograpHY (SCIAMACHY) (De Smedt et al., [2008;](#page-6-0) Wittrock et al., [2006\)](#page-9-0), Ozone Monitoring Instrument (OMI) (De Smedt et al., [2015\)](#page-6-0), GOME‐2A (De Smedt et al., [2012](#page-7-0)), and GOME‐2B (De Smedt et al., [2015](#page-6-0)).

Validation exercises with ground‐based Multi‐AXis Differential Optical Absorption Spectroscopy (MAX‐ DOAS) (Chan et al., [2020;](#page-6-0) De Smedt et al., [2021](#page-6-0); Ryan et al., [2023\)](#page-8-0) and Fourier-Transform InfraRed (FTIR) (Vigouroux et al., 2020) instruments show high temporal ($r = 0.71$ to 0.88) and spatial ($r = 0.85$ to 0.91) correlations, confirming the validity of TROPOMI HCHO product. Therefore, TROPOMI HCHO product has been widely used in constraining both biogenic and anthropogenic VOC emissions (Jin et al., [2023](#page-7-0); Zhao et al., [2022](#page-9-0); Zuo et al., [2023\)](#page-9-0) and the implication for ozone levels (Goldberg et al., [2022;](#page-7-0) Li et al., [2021\)](#page-7-0).

We select June 2018–May 2023 TROPOMI level 2 overpassing daily pixels with (a) quality assurance values greater than 0.5, (b) cloud fraction less than 0.3, (c) solar zenith angle less than 60° , and (d) snow-ice-free flag. We re-grid and oversample the resulted pixels onto the $0.5^\circ \times 0.5^\circ$ (~50 \times 50 km²) monthly grids, considering a tradeoff between ensuring sufficient pixels and reasonable computing speed, built on our Oversampling method (Pu et al., [2022;](#page-8-0) Sun et al., [2021](#page-9-0); Zhu et al., [2014](#page-9-0); Zhu, & Jacob et al., [2017](#page-9-0); Zhu, & Mickley et al., [2017;](#page-9-0) Zuo et al., [2023](#page-9-0)). The re-gridding method allows a precise allocation of pixels by area weight, and the oversampling method allows strict filtering criteria and increases the spatial signal-to-noise ratio by sacrificing temporal resolution. Further analysis is limited to grids where biogenic emissions are the main driver of HCHO columns (79% of the total continental grids), as determined from GEOS‐Chem sensitivity simulations to the sources from biogenic, anthropogenic, and biomass burning (Text S1 and Figure S1 in Supporting Information S1).

Figure 1. Global temperature (T-) dependency of TROPOMI HCHO columns. β_{TROPOMI} (unit: ln (10¹⁵ molecules cm⁻²) K^{-1}) is the TROPOMI HCHO T-dependency, fitted with $\Omega_{TROPOMI} = \exp(\alpha_{TROPOMI} + \beta_{TROPOMI} T + \varepsilon)$ at each $0.5^{\circ} \times 0.5^{\circ}$ grid, where $Ω_{TROPOM}$ is a vector of monthly oversampled TROPOMI HCHO columns and *T* is a vector of the monthly mean 2-m temperature (>285 K) correspond to $Ω$ _{TROPOMI} $α$ _{TROPOMI} is the fitting baseline and *ε* is the fitting residual. Only showing are biogenic-dominated grids (Figure S1 in Supporting Information S1) with significant $β_{TROPOMI}$ (*p*-value <0.05) and high determination coefficients (R^2 > 0.5, Figure S2 in Supporting Information S1). Low T-dependency regions (<0.02), such as the barren Sahara, are in gray and excluded from further analysis.

3. Global Temperature Dependency of TROPOMI HCHO Columns

We assign each oversampled HCHO column grid with a MERRA-2 surface temperature (T2M) according to its location and time. Here, we use air temperature as an indicator of leaf temperature, as they are positively and linearly correlated (Blonder & Michaletz, [2018;](#page-6-0) Fauset et al., [2018](#page-7-0); Still et al., [2022\)](#page-9-0). To quantify the Tdependency of HCHO columns, we apply the ordinary least squares regression to fit the exponential relationship between HCHO columns and temperature in each grid.

$$
\Omega = \exp(\alpha + \beta T + \varepsilon) \tag{1}
$$

where Ω (molecules cm⁻²) denotes a vector of monthly HCHO columns. *T* (K) is a vector of T2M with a baseline of 285 K. We discard the records with T2M lower than 285 K from the fitting, considering enzyme activity in BVOC productions is suppressed at low temperatures (Lehning et al., [1999](#page-7-0)), and isoprene emissions increase markedly at around 12°C (Goldstein et al., [1998](#page-7-0)). β , in the unit of ln (molecules cm⁻²) K⁻¹, quantifies the Tdependency of HCHO columns. α is the baseline after the logarithm, and ε is the fitting residuals. We focus on biogenic‐dominated grids (Figure S1 in Supporting Information S1) with significant *β* (*p*‐value <0.05) and temperature explaining the majority of HCHO variations ($R^2 > 0.5$) (Figure S2 in Supporting Information S1). Low T-dependency regions, such as barren and background areas, are filtered out from further analysis by a criterion of 0.02, the value over the Sahara Desert. Around 75% and 79% of the standardized residual from fitting grids over the research regions are considered normal and independent, respectively, indicating that the T‐ dependency fitting captures the main patterns and sources of variation. We admit that radiation flux (Zhang et al., [2019\)](#page-9-0) and leaf phenology (Surl et al., [2018](#page-9-0)) also jointly influence the HCHO variation with temperature in specific regions, such as Mid Amazon and North India.

As shown in Figure 1, the T-dependency of TROPOMI HCHO columns ($β$ _{TROPOMI}) has a more pronounced response over land. This highlights the dominant contribution of isoprene from terrestrial plants (Arnold et al., [2009;](#page-6-0) Guenther et al., [2006](#page-7-0), [2012](#page-7-0); Palmer & Shaw, [2005\)](#page-8-0), especially in Southeast US (McKinney et al., [2011](#page-8-0); Pressley et al., [2005;](#page-8-0) Stoy et al., [2021\)](#page-9-0) and Amazon (Alves et al., [2023;](#page-6-0) Wei et al., [2018\)](#page-9-0). *β*_{TROPOMI} decreases along the latitudes, with the maximum in the Amazon, tropical rainforests along the equator in Africa, and Southeast Asia, consistent with reported in situ isoprene emission hot spots (Jaars et al., [2016;](#page-7-0) Langford et al., [2022](#page-7-0); Misztal et al., [2011\)](#page-8-0). Such T-dependency appears to be a seasonal variation: summertime and rainy

Figure 2. Temperature (T-) dependency of TROPOMI ($β_{TROPOMI}$, blue), GEOS-Chem ($β_{GEOS-Chem}$, red), and perturbed GEOS-Chem (β[']GEOS-Chem, pink) HCHO columns across plant functional types (PFTs). The error bar is the standard deviation of the *β* for grids with the same PFT, demonstrating the variations in T‐dependency within the PFT. GEOS‐Chem results are corrected with TROPOMI averaging kernels (Text S2 in Supporting Information S1). The perturbed GEOS‐Chem incorporates β_{TROPOMI} into the temperature activity factor (γ'_T ; Equation [2\)](#page-4-0). Percentages quantify changes from $\beta_{\text{GEOS-Chem}}$ to β'_{GEOS}-Chem, relative to β_{TROPOMI}. PFTs are labeled in order: 1 Needleleaf Evergreen Temperate Tree, 2 Needleleaf Evergreen Boreal Tree, 4 Broadleaf Evergreen Tropical Tree, 5 Broadleaf Evergreen Temperate Tree, 6 Broadleaf Deciduous Tropical Tree, 7 Broadleaf Deciduous Temperate Tree, 8 Broadleaf Deciduous Boreal Tree, 10 Broadleaf Deciduous Temperate Shrub, 12 Arctic C3 Grass, 13 Cool C3 Grass, 14 Warm C4 Grass, and 15 Crop.

seasons have a notably higher T-dependency compared to wintertime and dry seasons (Figure S3 in Supporting Information S1).

We further examine mean β_{TROPOMI} across PFTs on land (Figure S4 in Supporting Information S1), by starting with the Broadleaf Evergreen Tropical Tree (PFT 4) dominated regions, such as Amazon, the equator of Africa, and Southeast Asia, where isoprene and monoterpene emissions contribute to 88% and 83% of the global total (Sindelarova et al., [2014\)](#page-8-0), respectively. We see from Figure 2 that β_{TROPOM} is the highest (0.12 \pm 0.07) for PFT 4, translating to that HCHO columns would double every 6.0 K \pm 4.1 K. This aligns with isoprene emissions variations measured in Amazon (Jardine et al., [2014](#page-7-0), [2016;](#page-7-0) Pétron et al., [2001\)](#page-8-0). *β*_{TROPOMI} is the lowest for the</sub> Broadleaf Deciduous Boreal Tree (doubling every 27.6 K \pm 4.8 K). As summarized in Table S1 of Supporting Information S1, the doubling temperature ranges from 9.2 to 30.8 K for other PFTs, enveloping values from previous observation‐based studies (Duncan et al., [2009](#page-7-0); Pressley et al., [2005\)](#page-8-0). The standard deviation is notable for some PFTs, such as Broadleaf Evergreen Trees (PFT 4 and 5) and Grasses (PFT 13 and 14), indicating a large variation in T-dependency within those PFT groups. Nevertheless, most of the T-dependency PFTs pairs are significantly different (Tukey pairwise post hoc test) except for 4 Broadleaf Deciduous species (PFT 7, 8, 10, and 12), acknowledging that the joint influence of latitudinal distribution on PFTs and T-dependency is nonnegligible, especially in the lower latitudes (Figure S5 in Supporting Information S1).

Figure [1](#page-2-0) also shows significant T-dependencies over the ocean, including off the coast around the Caribbean, the Gulf of Guinea, the ocean near Southeast India, the ocean near Western Australia, and coastal areas of the southeast US. We examine potential drivers of this by first correlating TROPOMI HCHO columns with the Greenhouse Gases Observing Satellite-2 (GOSAT-2) methane columns (Yokota et al., [2009\)](#page-9-0) (Figure S6 in Supporting Information S1). Over those regions with significant T-dependency, the correlation is either insignificant or negative, indicating other drivers rather than methane oxidation. Such T‐dependencies may be traced to seasonal changes of wind direction influencing the terrestrial outflow of long-lived VOCs (Gopikrishnan & Kuttippurath, [2021](#page-7-0); Wittrock et al., [2006\)](#page-9-0) and localized VOCs emitted by phytoplankton activities (Conte et al., [2020;](#page-6-0) Halsey et al., [2023\)](#page-7-0), admitting high uncertainties in oceanic emissions (Arnold et al., [2009](#page-6-0); Booge et al., [2016,](#page-6-0) [2018;](#page-6-0) Myriokefalitakis et al., [2010\)](#page-8-0).

GEOS-Chem HCHO Temperature Dependency

Figure 3. Global temperature (T-) dependency of GEOS-Chem HCHO columns. $\beta_{\text{GEOS-Chem}}$ (unit: ln (10¹⁵ molecules cm⁻²) K^{-1}) is defined as $\Omega_{\text{GEOS-Chem}} = \exp(\alpha_{\text{GEOS-Chem}} + \beta_{\text{GEOS-Chem}}T + \varepsilon)$ (Text S2 in Supporting Information S1). $\Omega_{\text{GEOS-Chem}}$ is the monthly GEOS-Chem HCHO column, corrected with the TROPOMI averaging kernels. *α*_{GEOS-Chem} is the baseline, and *T* denotes the 2-m temperature (>285 K). Only showing are biogenic-dominated grids with significant $β_{GEOS-Chem}$ (*p*value <0.05) and high dependency (R^2 > 0.5). Low T-dependency regions (<0.02; gray) are excluded from further analysis, consistent with β_{TROPOMI} in Figure [2.](#page-3-0)

4. Interpreting *β***TROPOMI With MEGAN Implemented in GEOS‐Chem**

To interpret how $β_{TROPOMI}$ varies spatially and across PFTs, we use the GEOS-Chem model (Text S1 in Supporting Information S1) that links HCHO columns with BVOC emissions driven by an array of environmental factors through the MEGAN module. GEOS‐Chem HCHO vertical profiles are sampled locally at TROPOMI over-passing time (13:00–14:00) and then corrected with TROPOMI averaging kernels (Text S2 in Supporting Information S1). As shown in Figure S7 of Supporting Information S1, GEOS‐Chem HCHO columns, in general, are linearly related to TROPOMI HCHO columns, supporting the interpretation of TROPOMI HCHO columns with GEOS-Chem, acknowledging that an overall overestimation and underestimation exist (Figure S7 in Supporting Information S1, panel c and d, respectively), which is not reflected by the T-dependency analysis.

We see from Figure 3 that the T-dependency of GEOS-Chem HCHO columns ($β$ _{GEOS-Chem}) consistently shows latitudinal variations and hot spots as reflected by $β_{TROPOM}$ (Figure [1\)](#page-2-0), but with an overestimation than TRO-POMI. Such overestimation is reflected in the overestimation of HCHO column variations (Figure S7 in Supporting Information S1, panel b). GEOS-Chem also overestimates the T-dependency in the Indian Ocean, off the coast of Western Australia, and around the Caribbean up to two times. This leads to the hypothesis that the ocean sources of HCHO may undermine the HCHO T-dependency offshore brought by seasonal winds (Gopikrishnan & Kuttippurath, [2021](#page-7-0)) since GEOS-Chem includes no ocean contribution. Additionally, *β*_{GEOS-Chem} coincides highly ($r = 0.87$) with β_{TROPOMI} on the PFT level (Figure [2\)](#page-3-0). On average, $\beta_{\text{GEOS-Chem}}$ is 16.0% higher than *β*TROPOMI, with discrepancy ranging from 6.1% (PFT 14, Warm C4 Grass) to 62.8% (PFT 8, Broadleaf Deciduous Boreal Tree).

To evaluate how BVOC emissions contribute to such discrepancy between $β$ _{GEOS}-Chem and $β$ _{TROPOMI}, we refer to the calculation of the temperature activity factor (γ_T) in the MEGAN module implemented in the GEOS-Chem model and perturb it using information from TROPOMI as follows:

$$
\gamma'_T = \gamma_T \times \sum_{i=1}^{15} \chi_i \exp(\Delta \beta_i (T - T_s)) \tag{2}
$$

where χ_i is the fractional area from PFT *i* coverage within the grid, $\Delta\beta_i = \beta_{\text{TROPOMI}, i} - \beta_{\text{GEOS-Chem}, i}$ corrects the discrepancy between $β_{TROPOMI}$ and $β_{GEOS-Chem}$ for PFT *i* (Table S2 in Supporting Information S1), *T* and *T_s* is the respective air temperature and standard temperature as described in the MEGAN model (Chen et al., [2023](#page-6-0); Guenther et al., [1993](#page-7-0); Zeng et al., [2023\)](#page-9-0) and set to be 303 K in GEOS-Chem. T-dependencies of HCHO columns from the perturbed GEOS-Chem simulation as denoted as $β'$ _{GEOS-Chem}.

We see from Figure [2](#page-3-0) that $β'$ _{GEOS}-Chem shows varying degrees of improvement compared with the default *β*_{GEOS}-Chem</sub>, with PFT 4 (Broadleaf Evergreen Tropical Tree) and PFT 14 (Warm C4 Grass) being the most efficient (72.9% and 90.8%, respectively). This indicates that BVOC emissions T-dependencies contribute mainly to HCHO columns T-dependencies at the lower and middle latitudes, where BVOC emissions are abundant. Our simulation shows that isoprene emission flux from PFT 4 under high temperatures, such as 315 K, is overestimated by at least 11% (\sim 1.2 \times 10⁻⁹ kg m⁻² s⁻¹) by the default GEOS-Chem. The result is consistent with the reported mismatch between MEGAN estimation and satellite-derived isoprene emissions (Barkley et al., [2012,](#page-6-0) [2013](#page-6-0)), recognizing that the degree of overestimation may vary depending on the versions of MEGAN and the landcover therein, and the meteorological information used in GEOS‐Chem.

The remaining discrepancy between $β_{GEOS-Chem}$ and $β_{TROPOMI}$ for other PFTs is likely due to the insensitivity of GEOS‐Chem HCHO columns to BVOC emissions. To verify this, we replace the default isoprene emissions with the measured fluxes (Seco et al., [2022\)](#page-8-0) in a boreal region (PFT 11, Broadleaf Deciduous Boreal Shrub). The sensitivity simulation shows that ∼80% of the isoprene emission variation only leads to ∼6.0% of the HCHO columns variation (Figure S8 in Supporting Information S1), indicating a weak response of HCHO columns to BVOC emissions. This reminds us that challenges exist when applying satellite HCHO columns as the proxy of BVOC emissions in regions with those PFTs, acknowledging other factors, such as Leaf Area Index (LAI), Aerosol Optical Depth (AOD), nitrogen oxides (NO*x*) level, and season transition, also complicate the relationship between HCHO columns and BVOC emissions (Alves et al., [2018](#page-6-0); Barkley et al., [2012,](#page-6-0) [2013;](#page-6-0) Marais et al., [2012;](#page-8-0) Strada et al., [2023\)](#page-9-0).

The perturbation improves T‐dependency in general but has both reductions and amplifications in the discrepancy of HCHO columns simulation. The discrepancy of HCHO columns mean value reduces in regions where GEOS‐ Chem has an overall overestimation compared with TROPOMI (mainly Southeast US and tropical regions), while amplifying in underestimation regions (mainly India and Australia) (Figure S9 in Supporting Information S1). On the PFT level, the perturbation improves the total HCHO columns by 1.8% for the Broadleaf Evergreen Tropical Tree, while exacerbating the discrepancy in the rest of the PFTs (0.47%–4.5%).

5. Conclusions

We have used TROPOMI satellite observations to investigate the temperature (T-) dependency of HCHO columns over the globe. The T-dependence of HCHO columns exhibits notable variations across Plant Functional Types (PFTs), with the highest dependency observed in the Broadleaf Evergreen Tropical Tree category. The GEOS-Chem model primarily interprets the T-dependency of HCHO columns at the PFT level, although there is a 16% discrepancy on average. We perturbed the temperature activity factor in MEGAN by introducing TROPOMI HCHO T-dependency information. The main drivers of such a discrepancy are the T-dependency of BVOC emissions for Broadleaf Evergreen Tropical Tree and Warm C4 Grass, and the insensitivity of HCHO columns to BVOC emissions for other PFTs.

We recommend a PFT-specific correction for temperature activity factor over tropical regions in the MEGAN model implemented in the GEOS-Chem, since BVOC T-dependency there interprets most of the HCHO Tdependency, and, although minor, has an improvement on HCHO columns simulation. The updates provide a correction accorded with satellite observation and improve the estimation of BVOC emission and HCHO columns, especially at high temperatures. We find that GEOS-Chem has an 11% (\sim 1.2 \times 10⁻⁹ kg m⁻² s⁻¹) of the isoprene emission overestimation compared with satellite observation for Broadleaf Evergreen Tropical Tree area at 315 K. The correction needs further investigation, such as additional model runs, including source attribution of HCHO T-dependency, since the improvement in this study is limited for those PFTs other than tropical regions.

Data Availability Statement

We gratefully acknowledge TROPOMI HCHO product (Copernicus Sentinel‐5P, [2020\)](#page-6-0), MERRA‐2 T2M product (Global Modeling and Assimilation Office, [2015\)](#page-7-0) implemented in GEOS‐Chem with the description access at [http://wiki.seas.harvard.edu/geos‐chem/index.php/MERRA‐2_implementation_details,](http://wiki.seas.harvard.edu/geos-chem/index.php/MERRA-2_implementation_details) GOSAT methane

Acknowledgments

This work is funded by the National Natural Science Foundation of China (42375090), National Key Research and Development Program of China (2023YFC3706205), Shenzhen Key Laboratory of Precision Measurement and Early Warning Technology for Urban Environmental Health Risks (ZDSYS20220606100604008), Guangdong Basic and Applied Basic Research Foundation (2021A1515110713), Guangdong University Research Project Science Team (2021KCXTD004), Major Talent Project of Guangdong Province (2021QN020924), Shenzhen Science and Technology Program (KQTD20210811090048025, JCYJ20220530115404009), and High level of special funds (G030290001). This work is supported by the Center for Computational Science and Engineering at Southern University of Science and Technology.

product (Japan Aerospace Exploration Agency, [2018](#page-7-0)). Oversampling code, data analysis scripts, and GEOS‐ Chem configuration files are available at the Zenodo repository (Li et al., [2024](#page-8-0)).

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