# AERONOMICA ACTA 

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A-N^{\circ} 21-1963
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## Solar radioflux and upper atmosphere temperatures by M. NICOLET

FOREWORD
"Solar Radioflux and UppereAtmosphere Temperature" is a paper presented at the COSPAR symposium in Warsaw (June 1963) and will be published in Journal of Geophysical Research, USA.

AVANT-PROPOS
"Solar Radioflux and Upper-Atmosphere Temperature" a été présenté au Symposium de COSPAR à Varsovie, juin 1963 et sera publié dans Journal of Geophysical Research, USA.

VOORWOORD
"Solar Radioflux and UppermAtmosphere Temperature", werd in juni 1963 voorgesteld op het symposium van COSPAR te Warschau en zal gepubliceerd worden in de Journal of Geophysical Research, USA.

VORWORT
"Solar Radioflux and Upper-Atmosphere Temperature" wurde wharend des COSPAR Symposium in Warschau (Juli 1963) vorgestellt und wird im Journal of Geophysical Research herausgegeben werden.

## Résumé.

Afin de déterminer la variation de la température de la haute atmosphère, au niveau de la thermopause, en fonction des conditions solaires, on a effectué une analyse génerale des flux radio-électriques solaires dans le domaine spectral de $10.000 \mathrm{Mc} / \mathrm{s}(3 \mathrm{~cm})$ a $1000 \mathrm{Mc} / \mathrm{s}$ ( 30 cm ).

La relation essentielle permettant de tenir compte du chauffage ultraviolet solaire peut etre trouvé dans le domaine spectral decimétrique en introduisant comme composante fondamentale du flux radioelectrique, une valeur moyenne couvrant une période de 27 jours. La comparaison effectuée entre les énergies émises à 8 cm et à 10 cm démontre qu'une relation linéaire existe au cours de tout le cycle solaire pour le flux émis à 8 cm tandis qu'une variation différente se présente pour le flux émis à 10 cm pour les conditions de faible activite solaire.

Les valeurs journalières du flux radio-électrique représentant ses fluctuations au cours d'une période de 27 jours interviennent dans la variation de la temperature a la thermopause avec un poids moindre que le flux moyen. L'ensemble des deux effets montre que la radiation ultraviolette solaire fournit le chauffage adéquat de la thermosphère et aucun effet corpusculaire n'est requis.

L'établissement des formules suivant le procédé utilisé par Jacchia permet de fournir a priori le tableau des températures et des flux solaires. Finalement, les temperatures moyennes en même temps que les températures maximales et minimales possibles ont éte déterminés pour les conditions de jour et de nuit de 1952 à 1962.

The correlation between the orbital deceleration of an artificial satellite and the solar radioflux in the decimeter range has led to relationships between the atmospheric density at various heights and the 10.7 and 20 cm radiofluxes. Considering that the density variation is due to a temperature variation, it mey be concluded that the solar ultraviolet radiation is responsible for the heating of the upper atmosphere and heat conduction is responsible for its cooling. A change in the ultraviolet radiation can be correlated with a change in the radioflux just as with the sunspot number, which is a well known index of solar activity.

Since all solar ultraviolet radiations do not vary with the same amplitude during a complete solar cycle, it appears logical to consider that a specific radiation in the radio range cannot be a perfectly reliable index. [1] Furthermore, various spurious correlations, or lack of correlations, may occur if the published values of the radiofluxes are subject to drifts of instrumental origin ${ }^{[2]}$. There is no way to account for a special behavior, for example at $5000 \mathrm{Mc} / \mathrm{s}$, when the entire spectrum between $1000 \mathrm{Mc} / \mathrm{s}$ and $10,000 \mathrm{Mc} / \mathrm{s}$ varies in exactly the same manner. Consequently, before carrying out an analysis of the various causes of the temperature variations of the thermopause, it is desirable to obtain reliable data on the variations of the solar radiofluxes.
2.- Variations of Solar Radiations.

Extensive solar observations have been conducted at a wavelength of $10.7 \mathrm{~cm}(2800 \mathrm{Mc} / \mathrm{s})$ (Covington [3]) at Ottawa by the Radio and Electrical Engineering Division of the National Research Counciawof Canada. The series of observations began in 1947, and the errors, which have been discussed by Medd and Covington [4] and Covington [5], should be of the order of $\pm 2$ percent, or not more $\pm 3$ percent in the monthly measurements. Intensive calibration measurements have been made at Ottawa with two radiometers calibrated fram two fixed points, one for the ambient temperature, and the otber at the temperature of the sky; taken to be $8^{\circ} \mathrm{K}$. A possible systematic error may be introduced, therefore, with an assumed constant sky temperature apparently
leading to a seasonal periodic variation proportional to the solar flux.
The routine observations at 8 cm have been carried out at Nagoya by the Institute of Atmospherics since November 1951 (Tanaka and Kakimuna ${ }^{[7]}$ ) with also an estimated accuracy of the order of $\pm 2$ percent. A comparison of the 10.7 cm and 8 cm fluxes made by Tanaka [8] during the period of sunspot minimum shows that the ratio is almost constant within $\pm 3$ percent. The ratio $10.7 \mathrm{~cm} / 8 \mathrm{~cm}$, which has a minimum in July and a maximum in January, indicates, however, the possibility of a periodic variation in the monthly average values which is dependent on the calibration. Other observations have also been conducted at Nagoya (Tanaka and Kakimuna [9]) at 3.2 cm since May 1956 and at 15 cm and 30 cm since the beginning of the International Geophysical Year (Tanaka and Kakimuna ${ }^{[10]}$ ). On that an attenuation due to the lower atmosphere may occur in the radiofluxes, particularly for the shortest wavelengths, systematic errors may be introduced in the published data since the transmission through unit atmosphere is subject to various variations. Such variations, generally proportional to the solar fluxes, may introduce periodic oscillations in the basic correlations between radiofluxes and upper-atmospheric parameters. These in addition to sporadic fluctuations which are not related to the solar emission.

The significance of radio observations for studying the physical conditions in the upper atmospheres of planets is well established since such observations provide indices of solar activity in addition to the zurich sunspot number (see for example, Kundu [11]). It is clear that a solar activity index can be chosen in the $1,000-10,000 \mathrm{Mc} / \mathrm{s}$ spectill range, but it appears that the spectral variation must be considered in the study of correlations with atmospheric phenomena. Since the radioflux is composed, without considering the burst emissions, of a quiet component (basic radio flux) and a slowly varying component, the variation of ultraviolet radiations which heat and ionize the planetary atmospheres must be compared along with that of the X-ray radiations. Friedman ${ }^{[12]}$ has clearly shown that the $X$-ray spectrum between 40 and 60 A varies by a factor of
approximately 7 during a solar cycle while the variation in the range 8 - 20A attains a value of about 50 and is several hundred in the 2 to 8 A wavelength range. Consequently the behavior in the $3-30 \mathrm{~cm}$ interval cannot be taken as representative of the solar X-ray flux or of the integral of the ultra-violet flux. The radioflux in the neighbourhood of 10 cm may be compared during a certain period of the solar cycle with an ionospheric index, such as the ionization of the E region, but cannot be expected to fit all such observations during an entire solar cycle. There is always a relationship between solar radio emissions in the $3-30 \mathrm{~cm}$ interval, X-ray and UV emission and atmospheric parameters but the correlations vary during the solar cycle.

Allen [13] and Covington and Harvey [14] have found that the quiet sun radioflux should have a significant variation with solar activity, which can reach a factor of about 2 , for the minimum-to-maximum variation in the 10 cm flux. The slowly varying component, which is associated with disturbed regions on the solar disk and is closely correlated with sunspot numbers, leads, in the same spectral region, to a minimum-to-maximum variation of at least a factor of 4. A recent analysis by Das Gupta and Basu [15] shows that, during the International Geophysical Year (July 1957 to December 1958, maximum of solar activity), there is a close correlation between the sunspot number and daily flux between 3 and 30 cm , and furthermore, that the maximum variation occurs near $8-10 \mathrm{~cm}$.

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3.- Solar Cycle Variations.
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Since the solar radio-emissions in the range of wavelengths from 3 cm to 30 cm show similar variations with a so-called 27 day cycle, it is of interest to plot the 27 -day average value of the Zurich relative sunspot number and of the solar radio flux at 10 cm from 1957 till 1962. The plot (Fig. 1) of the Wolf number ${ }^{(*)} \overline{\mathrm{R}}_{27}$ with the radioflux ${ }^{(* *)} \overline{\mathrm{S}}_{27}(10.7 \mathrm{~cm})$
(*) 27-day geomagnetic period (Bartels) deduced from Zurich Observatory data. (**) deduced from Ottawa data in units of $10^{-22}$ watts/m ${ }^{2} /$ cyc1e/sec.


Fig. 1.- Solar radioflux at 10.7 cm versus relative sunspot number between 1951 and —— 1962 when 27 -day mean values are used.
shows that, for $\overline{\mathrm{R}} \geq 50$, there is a linear relationship between these quantities which, with an exror of $\pm 10 \%$, takes into account the variations of the radioflux for values greater than $S=100$ units, i.e.

$$
\overline{\mathrm{s}}_{27}(10.7 \mathrm{~cm})=50+0.967 \overline{\mathrm{R}}_{27}
$$

But for $\overline{\mathrm{S}}<100$ or $\overline{\mathrm{R}}<50$, a linear variation would be:

$$
\overline{\mathrm{S}}_{27}(10.7 \mathrm{~cm})=68+0.607 \overline{\mathrm{R}}_{27}
$$

Thus, the close association of the 10 cm radiation with sunspots cannot be expressed during a whole cycle with the same linear relationship. Furthermore, the dispersion of the daily values is larger than the 27-day mean values. Fig. 2, which shows a plot of daily values with a variation reaching $\pm 20$ percent, indicates how such a solar index may lead to important errors in deducing any empirical correlation with atmospheric parameters.

To determine the differences between two radio emissions we can compare, first, the 27 -day mean values at two wavelengths, such as 8 cm and 10.7 cm for which data are available since 1952. Even if the absolute measurements of radiofluxes are estimated with an error of $\pm 7$ percent, and can deviate by such a percentage from a ratio of unity after normaiisation, the general variation during a whole solar cycle should be significant [16]. Fig. 3 shows the gradual difference in the ratio of the 27 -day mean values of solar fluxes at 10.7 cm and 8 cm . Near 200 units, both fluxes are identical, i.e. $\overline{\mathrm{s}}_{27}(10.7)=\overline{\mathrm{s}}_{27}(8)$. For the lower values $\overline{\mathrm{s}}_{27}(8)>\overline{\mathrm{s}}_{27}$ (10.7) while the ratio is reversed, $\overline{\mathrm{S}}_{27}(10.7)>\overline{\mathrm{S}}_{27}$ (8) for the higher values, it is clear that the ratio $\overline{\mathrm{S}}_{27}(10.7) / \overline{\mathrm{S}}_{27}$ (8) varies during a solar cycle, even if the emissions at these two wavelengths originate from almost identical levels in the solar atmosphere.

If we compare the variation of fluxes for two years of high and low solar activity (Fig. 4 and 5, 1958 and 1961), the annual mean values are :


Fig. 2.- Solar radioflux at 10.7 cm versus relative sunspot number for daily values in 1958 near sunspot maximum.


Fig. 3.- Solar radiofluxes at 10.7 cm and 8 cm . 27-day mean values between 1952 ——— and 1962.


Fig. 4.- Distribution of daily values of solar radiofluxes at 10.7 cm and 8 cm in 1958 near sunspot maximum.


$$
\frac{\bar{S}_{1958}(8)}{\bar{S}_{1958}(10.7)}=\frac{224.8}{230.6}=0.975
$$

and

$$
\frac{\bar{s}_{1961}(8)}{\bar{s}_{1961}(10.7)}=\frac{114.9}{113.9}=1.09
$$

There is, therefore, a reversal in the ratio of the 8 cm and 10.7 cm radiofluxes when there is a variation of a factor of about two in the absolute values of the fluxes. Furthermore, the maximum amplitude of the variation of daily fluxes, compared with the annual mean value, is, in both cases, of the order of $\pm 30 \%$. At sunspot maximum (1958, Fig. 4) the fluctuations of the ratio $\left[S(8) / \bar{S}_{1958}(8)\right] /\left[S(10.7) / \bar{S}_{1958}(10.7)\right]$ are generally less than $\pm 5$ percent and always less than $\pm 10$ percent. However, the same ratio in 1961 (Fig. 5)is represented by scattered values reaching $\pm 15 \%$ but with systematic differences when the ratio is greater than or less than unity. In fact, the amplitude of the variation at 8 cm appears to be less than that at 10 cm . Thus, it can be concluded that the systematic differences occuring in the annual average, 27-day average and daily values of fluxes measured at 10.7 cm (Ottawa) and at 8 cm (Nagoya) are a clear indication of real variations in the spectrum of the solar radiofluxes between $2800 \mathrm{Mc} / \mathrm{s}$ and $3750 \mathrm{Mc} / \mathrm{s}$.
4.- Spectrum from $1,000 \mathrm{Mc} / \mathrm{s}$ to $10,000 \mathrm{Mc} / \mathrm{s}$.

A comparison between the fluxes at 15 cm and 30 cm , which are also measured at Nagoya, and the 10.7 cm flux does not show systematic differences such as those whsurt are noted between 8 cm and 10.7 cm . From Fig. 6, it appears that the ratios have the following median values :

$$
\overline{\mathrm{s}}_{27}(15) / \overline{\mathrm{s}}_{27}(10.7)=0.72 \pm 0.08
$$



Fig. $6 .=$ Variation of ratios of 27 -day mean values of solar radio fluxes at 8 cm , -15 cm and 30 cm and radioflux at 10.7 cm .
and

$$
\bar{s}_{27}(30) / \bar{s}_{27}(10.7)=0.55 \pm 0.07
$$

In other words, the fluctuations, of the order of $\pm 10$ percent, are of sufficiently small magnitude to be explained by experimental errors without showing real spectral differences. In any event, such a variation of the ratios cannot be indicative of a systematic difference associated with the solar cycle as has been found for 8 cm . In fact, the ratio of two radiofluxes really becomes sensitive to the solar cycle when the radiation is at a wavelength shorter than 10 cm .

The ratios $\bar{S}_{27}(3.2) / \bar{S}_{27}(8)$ and $\overline{\mathrm{S}}_{27}(3.2) / \overline{\mathrm{S}}_{27}$ (10.7) are considered in Fig. 7. The two curves lead to the median value of $\bar{S}_{3.2} / \bar{S}_{10.7}=1.5$ for $\bar{S}_{27}(10.7)>200$ with a maximum fluctuation of $\pm 10$ percent. The ratio increases with a decrease a solar activity. It is so sensitive to solar activity that its fluctuations follow the variations of the solar flux. Peaks in the ratio correspond to decreasing fluxes, and the minimum value of the ratio, namely 1.3 , is real and is associated with the highest activity of the sun in 1957. Thus, the variation of radiofluxes with solar activity in the spectral range $1,000 \mathrm{Mc} / \mathrm{s}$ $10,000 \mathrm{Mc} / \mathrm{s}$ is always expressed as follows :


It is evident that an adequate solar radio index may be chosen at 10 cm or at shorter wavelengths. However, a more complete analysis must be made before adopting a general law for the entire solar cycle. Fig. 8 shows how the ratio $\bar{S}_{27}(8) / \bar{S}_{27}(10.7)$ varies between 1952 and 1962. The ratio reaches a maximum at the minimum of the solar cycle with a median value of $1.15 \pm 5$ percent and decreasing to a median value of $0.98 \pm 5$ percent at the maximum of the solar cycle, i.e. a difference


Eig. 7.- Effect of solar activity on the spectral variation between 10 cm and 3 cm .


Fig. 8.- Variation of the ratio of 27 day mean values of solar radiofluxes at 8 cm and 10.7 cm from 1952 to 1962.
reaching 20 percent for the minimum-to-maximum variation of a solar cycle. It is clear, therefore, that the absolute fluxes must change as indicated by this ratio.

To give an idea of the amplitude of the 27 -day variation of the slowly varying component, Figs 9 and 10 show the ratio $\mathrm{S}_{\mathrm{Max}} / \mathrm{S}_{27}$ and $\mathrm{S}_{\mathrm{Min}} / \mathrm{S}_{27}$, representing the highest and lowest values of the daily fluxes for each 27-day period. No systematic difference can be immediately detected with the solar cycle for 10 cm and 8 cm . It is clear that the amplitude of the 27 -day cycle at 3 cm is always less than that at 8 cm . As far as 15 cm is concerned, it appears that for $S(8) \geqslant 200$ units, i.e. for high solar activity, $\Delta S(15 \mathrm{~cm}) \leqslant \Delta S(8 \mathrm{~cm})$ while $\Delta S(15 \mathrm{~cm})>\Delta S(8 \mathrm{~cm})$ at low solar activity, i.e. when $S(8)<120$. The same law seems also to be adequate for 30 cm , since $\Delta S(30 \mathrm{~cm})<\Delta S(8 \mathrm{~cm})$ when $S(8)>150$ and $\Delta S(30 \mathrm{~cm}) \geq \Delta S(8 \mathrm{~cm})$ when $S(8)<120$. Such striking features exhibited by the 27 -day cycle between 8 cm and 30 cm indicate that its maximum-tominimum amplitude as detected by the daily mean fluxes, cannot be subject to simple laws related to the frequency. A systematic variation is found at high solar activity during the International Geophysical Year period (for example Das Gupta and Basu [15]), but is completely different in 1961 and 1962 as shown in Fig. 9.

Thus, a complete analysis of the solar radiofluxes should be made at wavelengths shorter than 10 cm since no clear conclusion (systematic errors or spectral differences) can be reached for the spectral range of wavelengths greater than 10 cm .
5.- Atmospheric Temperatures and Solar Fluxes.

Since the increase of the atmospheric scale height with altitude is due to a normal decrease of the molecular mass and not to an increase of the temperature [1], and the thermopause temperature is related to the ultraviolet heating and conduction cooling, a relationship can be found between the solar radioflux and the thermopause temperature. Jacchia [17],


Fig. 9.- Minimum-to-maximum variations of solar radiofluxes at 8 cm and 10.7 cm for a 27-day period between 1952 and 1962.


Fig. 10. - Minimum-to-maximum variations at $3.2 \mathrm{~cm}, 15 \mathrm{~cm}$, 30 cm and 8 cm between 1957 and 1962.
after adopting Nicolet's atmospheric model [18], deduced an equation having the following form :

$$
\begin{equation*}
T=T_{0}+a S \tag{1}
\end{equation*}
$$

In which the temperature of the thermopause, $T$, is related to the solar flux, $S$, with $d T / d S=a$ and $T_{o}$ a constant. Linear relationships such as (1) have also been adopted by Harris and Priester [19] and Paetzold [20].

From the preceding discussion it is clear that the adoption of a radio index at different wavelengths will modify the value of the coefficient a in equation (1). Further, these different coefficients will be obtained depending on the wavelength chosen.

Let us again consider the 5 monochromatic radiofluxes at 3.2 cm , $\therefore 8 \mathrm{~cm}, 10.7 \mathrm{~cm}, .15 \cdots \mathrm{~cm}$ and 30 cm for which the analyses made in the preceding paragraphs have shown the general behavior for a complete solar cycle. The following ratios were determined :
(i) Ratio: of the maximum 27-day mean value reached in 1957 (period $n^{\circ} 1701$ or 1703) and the minimum 27 -day mean value reached in 1954 (periods near. $n^{\circ}$ 1655) or, when no continuous observations are available, individual values during $n^{\circ} 1766$ in August 1962.
(ii) Ratio of the minimum daily fluxes reached in 1957 (periods between $\mathrm{n}^{\circ} 1695$ and $\mathrm{n}^{\circ} 1715$ ) and the minimum 27 -day mean value.
(iii)Ratio of the maximum of maximum daily fluxes reached in 1957 (period $n^{\circ}$ 1703) and the minimum 27 -day mean value.

The first.ratio, $R_{i}$, will give the maximum amplitude during a"solar cycle of the 27 -day mean value of the radioflux. The second ratio, $R_{i 1}$, corresponds to the maximum possible variation of the quiet sun during a. solar cycle and the third ratio, $R_{\text {iii }}$, corresponds to the minimam-tomaximum variation which can be reached. The vailous ratios aie sluwn in Table-I: A factor of about 6, which represents the minimum-to-maximum variation $R_{\text {iii }}$ near $2000 \mathrm{Mc} / \mathrm{s}$, corresponds to an extreme case. It is still,

TABLE $I_{\text {. - }}$ Ratios of various maxima of solar fluxes to the minimum values for

| $\mathrm{Mc} / \mathrm{s}$ | a quiet sun at minimum solar activity. |  |  | 2000 | 1000 |
| :---: | :---: | :---: | :---: | :---: | :---: |
|  | 9400 | 3750 | 2800 |  |  |
| $\lambda(\mathrm{cm})$ | 3.2 | 8 | 10.7 | 15 | 30 |
| $\mathrm{R}_{\boldsymbol{i}}$ | 1.6 | 3.7 | 4.2 | 4.5 | 4.5 |
| $\mathrm{R}_{\text {i }}$ | 1.2 | 2.2 | 2.5 | 2.8 | 3.0 |
| $\mathrm{R}_{\text {iii }}$ | 2.0 | 5 | 5.5 | 5.9 | 5.4 |

however, less than the variation of the X-ray flux (Friedman [12]), absorbed in the ionospheric $E$ layer. In fact, a factor of 7 in the $X$-ray emission must be compared with the maximum ratio $R_{i}$ of 4.5 representing the 27 -day meam ratio. The ratio $R_{i 1}$, which varies from 1.2 at 3.2 cm to 3 at 30 cm , shows how difficult it is to determine the amplitude of the quiet sun variation which should be related, first of all, to the ultraviolet flux associated with the solar heating of the upper atmosphere.

Nevertheless, since the ratio $R_{i i i} / R_{i}$ is of the order of 1.3 and the ratio $R_{i i} / R_{i}$ is of the order of 0.6 between 8 cm and 15 cm , equation (1) is used in a different form in which a basic component corresponds to the mean value $\bar{S}_{27}$ for a 27 -day period and a varying component $\Delta$ s represents the oscillation during such a period :

$$
\begin{equation*}
T=T_{0}+a \bar{S}_{27}+b \Delta S \tag{2}
\end{equation*}
$$

where

$$
\begin{equation*}
\Delta S=S-\bar{S}_{27} \tag{3}
\end{equation*}
$$

or

$$
\begin{equation*}
T=T_{0}+(a-b) \bar{\delta}_{27}+b \delta \tag{4}
\end{equation*}
$$

The coefficients $a$ and $b$ will have different values. $a$ will be greater than $b$ since it must take into account the basic component of the solar
flux primarily associated with the ultraviolet flux. The factor b represents the 27-day oscillation of a radioflux which is less important for the ultraviolet variation. In other words, considering the ratios $R_{i}$ and $R_{i i i}$, the term a $\bar{S}_{27}$ in (4) is associated with the variation of $R_{i}$ during a whole solar cycle and the term $b \Delta S$ corresponds to the fluctuations ( $R_{i i i}=R_{i}$ ) during a 27-day period.

In order to determine the variations for an entire solar cycle, one may compare the ratios of $R_{i}$ for $3.2 \mathrm{~cm}, 8 \mathrm{~cm}$ and 10.7 cm since the variation for the ultraviolet heating must be associated with this spectral range. The variation of $\bar{S}_{27}(10.7 \mathrm{~cm})$ is plotted against the ratio $\overline{\mathrm{S}}_{27}(8) /$ $\bar{S}_{2 j}(10.7)$ in Fig. 11 with the unit ratio at $S(10.7 \mathrm{~cm})=250$. A smoothed relation is drawn and indicates that the slope is steeper for lower activity. The correction factor reaches 20 percent at sunspot minimum and, it should be mentioned that, one would not expect a severe departure for the smooth curve since the individual ratios are affected with a maximum error of $\pm 5$ percent, i.e. not more than the possible errors in the absolute values. Fig. 12 represents the characteristic difference in the 27 -day mean fluxes of 3.2 cm and 10.7 cm . The normalisation of the ratio $\overline{\mathrm{S}}_{27}(3.2) / \overline{\mathrm{S}}_{27}(10.7)$ is made for $250<\overline{\mathrm{S}}_{27}(10.7)<300$ units. There is a significant change with the solar cycle, reaching a factor of 2.6 . The dashed-line curve corresponding to the ratio $\overline{\mathrm{S}}_{27}(8) / \overline{\mathrm{S}}_{27}(10.7)$, compared with the 3.2 cm curve, shows how the spectrum varies with solar activity between $3,000 \mathrm{Mc} / \mathrm{s}$ and $10,000 \mathrm{Mc} / \mathrm{s}$. The conclusion arrived at from such an analysis is that the variations of the ultraviolet radiation can be correlated with a radio flux in the range $10 \mathrm{~cm}-30 \mathrm{~cm}$ during the maximum of the solar cycle. However, the relative change of flux is systematically less for higher frequencies, and can be regarded as being more representative of the change of the ultraviolet spectrum related to the heating of an upper atmosphere. Since there are no measurements available between $10,000 \mathrm{Mc} / \mathrm{s}$ and $3,750 \mathrm{Mc} / \mathrm{s}$, a complete analysis can only be made at 8 cm and 10.7 cm . Jacchia $[21]$, using a smoothed value $\bar{S}(10.7 \mathrm{~cm})$ such as :

$\stackrel{\sim}{i}$

Fig. $11 .{ }^{-}$Ratio of solar radiofluxes at 8 cm and 10.7 cm .


Fig. 12.- Comparison between solar radiofluxes at 3.2 cm and at 8 cm with 10.7 cm .

$$
\begin{equation*}
\bar{T}=T(S=200)+2^{\circ} 5 \bar{S}(10.7 \mathrm{~cm}), \tag{5}
\end{equation*}
$$

has determined average night-time and day-time temperatures corresponding to an average value for the diurnal bulge. The result is that, except for the lowest values of $\overline{\mathrm{S}}$, a linear correlation with an average slope of $4^{\circ} 5$ between night-time temperatures and $\overline{\mathrm{S}}(10.7 \mathrm{~cm})$ may be found. The same linear relationship is also used by Harris and Priester [19], i.e.

$$
\begin{equation*}
\overline{\mathrm{T}}_{\mathrm{night}}=275^{\circ} \mathrm{K}+4^{\circ} 5 \overline{\mathrm{~S}}(107 \mathrm{~cm}) \tag{6}
\end{equation*}
$$

Adopting the same formula for 8 cm , we see (Fig. 11) that equation (8) below may be used if we adopt for the average slope of the variation of the night-time temperature :

$$
\begin{equation*}
\frac{\mathrm{dT} \text { (night) }}{\mathrm{dS}_{27}(8 \mathrm{~cm})}=4^{\circ} 6 \tag{7}
\end{equation*}
$$

The use of the 8 cm flux as a solar index leads to a basic correlation between the temperature and the flux at 10.7 cm which is no longer linear. Taking into account the differences of $\pm 5$ percent occuring between 8 cm and 10 cm fluxes we use, instead of (6),

$$
\begin{equation*}
\vec{T}_{\text {night }}=280^{\circ} \mathrm{K}+\left(4^{\circ} 6 \pm 0.25\right) \overrightarrow{\mathrm{S}}_{27}(8 \mathrm{~cm}) \tag{8}
\end{equation*}
$$

The variation of $\pm 5$ percent in the second term of (8) shows that differences of at least $\pm 0.25 \mathrm{~S}$ should be expected in any determination of a correlation between the temperature and the 27 -day mean value of any observed radioflux.

In order to fit the temperatures of the day-time bulge we use
Jacchia's relation

$$
\begin{equation*}
\bar{T}_{\text {day }}=\bar{T}_{\text {night }} \times 1.33 \tag{9}
\end{equation*}
$$

leading to $\bar{T}_{\text {day }}=370^{\circ} \mathrm{K}+(6.1 \pm 0.3) \overline{\mathrm{S}}_{27}(8)$

The maximum day time temperature may require a factor a little greater than 1.33 , but the following expression, equivalent to that used by Harris and Priester [19],

$$
\begin{equation*}
\bar{T}_{\text {dayotime maximum }}=\bar{T}_{\text {night }} \times 1.35+\bar{S}(10.7 \mathrm{~cm}) \tag{11}
\end{equation*}
$$

is not adequate since it leads to diurnal amplitudes which are too large and, furthermore, to systematic differences, depending on the level of solar activity, with temperatures which are too high.

Thus, equations (8) and (10) have been used to determine the diurnal oscillation with the variation of $\overline{\mathrm{s}}_{27}(8 \mathrm{~cm})$. In Fig. 13 the temperature variation is plotted against the usual 10.7 cm solar flux utilizing a re-conversion from 8 cm which shows the non-linearity of the relationships. The minimum-to-maximum variafion of the average night-time temperature during a solar cycle is from about $650^{\circ} \mathrm{K}$ for the completely quiet sun to $1600^{\circ} \mathrm{K}$ for the highest solar activity. The same variation for day-time temperatures corresponds to the range of $850^{\circ} \mathrm{K}$ to $2100^{\circ} \mathrm{K}$.

In order to use the 27 -day mean values of the 10.7 cm radioflux at low solar activity the following formulas may be adopted instead of (8) and (10) :

$$
\begin{equation*}
\overline{\mathrm{T}}_{\text {night }}=380^{\circ} \mathrm{K}+\left(4^{\circ} \pm 0.25\right) \stackrel{\rightharpoonup}{\mathrm{S}}_{27}(10.7 \mathrm{~cm}) \tag{8a}
\end{equation*}
$$

and

$$
\begin{equation*}
\overline{\mathrm{T}}_{\text {day }}=500^{\circ} \mathrm{K}+\left(5^{\mathrm{c}} 3 \pm 0.3\right) \overline{\mathrm{S}}_{27}(10.7 \mathrm{~cm}) \tag{10a}
\end{equation*}
$$

if $\mathrm{S}_{27}(10.7 \mathrm{~cm})<200$ units.
Jacchia and Slowley [22] have made a precise analysis of the 27-day oscillation deduced from the drag of several satellites. The temperature variation may be expressed by a relation such as


Fig. 13.- Average night-time and daytime temperatures of the thermopause for a complete solar cycle.

$$
\begin{equation*}
T=\bar{T}+b \Delta S \tag{12}
\end{equation*}
$$

with values of the coefficient $b$ which may be of the order of 2.5 ; extreme values of the order of $2^{\circ}$ and $3^{\circ}$ could be used for night and day-time conditions. However, it is not yet certain that there is a real difference between dark and sunlit atmospheres since Jacchia and Slowley considered the possibility of $a$ constant value of $b=2.2$.

Thus the final formula may be written, according to (2), (4), (8) and (12), and neglecting the possible error of $\pm 5$ percent,

$$
\begin{equation*}
T_{\text {night }}=280^{\circ} \mathrm{K}+4: 6 \overline{\mathrm{~S}}_{27}(8 \mathrm{~cm})+2: 5 \Delta \mathrm{~S} \tag{13}
\end{equation*}
$$

or

$$
\begin{equation*}
T_{\text {night }}=280^{\circ} \mathrm{K}+2.1 \overline{\mathrm{~s}}_{27}(8 \mathrm{~cm})+2.5 \Delta \mathrm{~S}(8 \mathrm{~cm}) \tag{14}
\end{equation*}
$$

This relation holds, as a general relation, between the solar flux at 8 cm and atmospheric temperature over a complete solar cycle and including the 27-day oscillation. Consequently, there is no need for an explanation of atmospheric heating involving a solar wind or any other corpuscular effect. In fact, the term involving $\overline{\mathrm{S}}_{27}$ is the basic component for the ultraviolet heating while the other term incorporating $\Delta S$ involves, with a smaller weight, the effect of the 27 -day oscillation. However, when both $S$ and $\bar{S}$ are used in the same formula the coefficients are of the same magnitude as deduced from the observations.

There are other effects which have also been introduced, such as the semiannual or annual variations detected by Paetzold [20], for which an origin has been suggested as an effect of a solar or an interstellar wind; respectively. The fact, however, that a solar index such as 8 cm instead of 10.7 cm flux has removed an essential difficulty, such as a lack of a linear correlation at low solar activity and also a systematic drift in the atmospheric temperature from 1958 to 1962, leads to other possible explanations for such effects which may be detected only after
several reductions of observational data. Several factors must be kept in mind : (i) Any atmospheric model may lead to systematic differences in the temperature of the thermopause of at least $\pm 50^{\circ} \mathrm{K}$. (ii) The boundary conditions in the thermosphere cannot be kept constant for an entire year. (iii) A steady state of diffusion conditions cannot be assumed in a detailed study of the diurnal variation. (iv) The absolute value of any solar radioflux may lead to an undetectable error of $\pm 5$ percent.
(v) Depending on the type of calibration, systematic drifts due to seasonal atmospheric effects can occur. Thus, fluctuations may occur which lead to possible singular behavior of certain parameters. In spite of these difficulties, the solar radioflux at 8 cm establishes a simple pattern which is suitable for representing the average temperature conditions of the earth's upper atmosphere at various stages of the solar cycle. This requires a heating by ultraviolet radiation and a cooling by heat conduction without postulating any essential corpuscular process. The later would lead to an absorbed energy of not less than $0.2 \mathrm{erg} \mathrm{cm}^{-2} \mathrm{sec}^{-1}$ while the normal atmospheric conditions show that it cannot exceed $0.05 \mathrm{erg} \mathrm{cm}^{-2} \mathrm{sec}^{-1}$. These influences must be found in the auroral zone, and are important during magnetic storms through hydromagnetic heating.
6.- Temperature Variations during a Solar Cycle.

Since the ratio $\overline{\mathrm{S}}_{27}(8) / \overline{\mathrm{S}}_{27}(10.7)$ varies with solar activity and there is no general relation which is applicable, numerical values must be used. In order to show how the differences may be dealt with a calculation has been made for average night-time and dayotime temperatures corresponding to the 27 -day mean values of solar radiofluxes at 8 cm and 10.7 cm by using formulas (6) and (8). Fig. 14 shows how the night-time temperatures deduced from the 10.7 cm radioflux would be systematically too low when the temperature is less than $1250^{\circ} \mathrm{K}$, the difference increases with the decrease of the solar activity. Fig. 15 gives the day-time temperatures from $850^{\circ} \mathrm{K}$ to $2100^{\circ} \mathrm{K}$ with a systematic difference below $1750^{\circ} \mathrm{K}$. It should be remarked that the scatter of the various points represents differences in the measured solar


Fig. 14.- Differences in the determination of the thermopause temperatures —— according to rat\&o $\mathrm{S}_{27}(8 \mathrm{~cm}) / \mathrm{S}_{27}(10.7 \mathrm{~cm})$ used as found in Fig. 11 。


Fig. 15.- Differences in the determination of the thermopause temperatures according - to ratio $\bar{S}_{27}(8 \mathrm{~cm}) / \bar{S}_{27}(10.7 \mathrm{~cm})$ used as found in Fig. 11.
fluxes. These must be interpreted as fluctuations representing possible errors when use is made of a solar radioindex which is measured with very great precision.

The possible error of about $\pm 0.25 \mathrm{~S}$ was not included in the following Table II, which gives, for various temperatures of the thermopause, the corresponding night-time and day-time radiofluxes. These fluxes correspond to 27 -day mean values at 8 cm and 10.7 cm . Table III gives all 27-day means of the temperature values from 1952 to 1962 , including the minimum-to-maximum values for each period. They are based on the linear relationships deduced in the preceding paragraph by usi:ng the 8 cm radioflux. Figs 16 and 17 yield a clear representation uf the average temperatures at the thermopause level for a 27 -day period. The vertical lines indicate the extreme minimum and maximum temperatures which have been reached during each period. The minimum temperatures are $650^{\circ} \mathrm{K}$ and $850^{\circ} \mathrm{K}$ for day and night-time temperatures ; respectively. The highest 27-day mean value for night-time conditions was reached in 1957 and is of the order of $1600^{\circ} \mathrm{K}$. The maximum night-time value, of the order of $1800^{\circ} \mathrm{K}$ was, however, attained during that pericd. The highest 27 -day mean value for day-time conditions is of the order of $2150^{\circ} \mathrm{K}$ while the maximum day-time value reached $2400^{\circ} \mathrm{K}$.

On conclusion, it should be noted that the effect of magnetic storms has not been introduced and shouid be incorporated in the determinations of temperature by using Jacchia's formula along with Ap indices.

TABLE II. - Thermopause temperatures and solar radiofluxes (*)

| $\begin{aligned} & \text { Temperature } \\ & \left({ }^{\circ} \mathrm{K}\right) \end{aligned}$ | NIGHT-TIME |  | DAY-TIME |  |
| :---: | :---: | :---: | :---: | :---: |
|  | $\bar{S}_{27}(8 \mathrm{~cm})$ | $\bar{S}_{27}(10.7 \mathrm{~cm})$ | $\bar{S}_{27}(8 \mathrm{~cm})$ | $\bar{S}_{27}(10.7 \mathrm{~cm})$ |
| 650 | 80 | 68 |  |  |
| 700 | 91 | 80 |  |  |
| 750 | 102 | 93 |  |  |
| 800 | 113 | 105 |  |  |
| 850 | 124 | 118 | 79 | 68 |
| 900 | 135 | 130 | 87 | 76 |
| 950 | 146 | 143 | 95 | 85 |
| 1000 | 157 | 155 | 104 | 94 |
| 1050 | 167 | 168 | 112 | 104 |
| 1100 | 178 | 180 | 120 | 113 |
| 1150 | 189 | 193 | 128 | 122 |
| 1200 | 200 | 205 | 136 | 132 |
| 1250 | 211 | 217 | 145 | 141 |
| 1300 | 222 | 228 | 153 | 151 |
| 1350 | 233 | 239 | 161 | 160 |
| 1400 | 244 | 250 | 169 | 169 |
| 1450 | 255 | 261 | 177 | 178 |
| 1500 | 265 | 272 | 186 | 188 |
| 1550 | 276 | 282 | 194 | 197 |
| 1600 | 287 | 293 | 202 | 206 |
| 1650 | 298 | 304 | 210 | 215 |
| 1700 | 309 | 315 | 218 | 223 |
| 1750 |  |  | 227 | 232 |
| 1800 |  |  | 235 | 240 |
| 1850 |  |  | 243 | 248 |
| 1900 |  |  | 251 | 257 |

(*) Solar radiofluxes given with precision of $\pm 0.25$ S corresponding to differences in observational data.

|  | NIGH | -TIME |  | -TIME |
| :---: | :---: | :---: | :---: | :---: |
| Temperature ( ${ }^{\circ} \mathrm{K}$ ) | $\overline{\mathrm{s}}_{27}(8 \mathrm{~cm})$ | $\overline{\mathrm{s}}_{27}(10.7 \mathrm{~cm})$ | $\bar{S}_{27}(8 \mathrm{~cm})$ | $\overline{\mathrm{s}}_{27}(10.7 \mathrm{~cm})$ |
| 1950 |  |  | 259 | 265 |
| 2000 |  |  | 268 | 273 |
| 2050 |  |  | 276 | 282 |
| 2100 |  |  | 284 | 290 |
| 2150 |  |  | 292 | 298 |
| 2200 |  |  | 300 | 306 |

(*) Solar radiofluxes given with precision of $\pm 0.25$ S corresponding to differences in observational data.

TABLE III.- Day-time and night-time temperatures at the thermopause.

| Rotation | $\begin{gathered} \text { Date } \\ ( \pm 13 \text { days }) \end{gathered}$ |  |  | Day-time |  |  | Night-time |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  |  | Max. | Average | Min. | Max. | Average | Min. |
| 1623 | 1952 | January | 17 | 1092 | 1045 | 1013 | 821 | 786 | 762 |
| 1624 |  | February | 13 | 1043 | 990 | 951 | 784 | 744 | 715 |
| 1625 |  | March | 11 | 942 | 910 | 895 | 708 | 684 | 673 |
| 1626 |  | April | 7 | 978 | 966 | 924 | 735 | 726 | 695 |
| 1627 |  | May | 4 | 986 | 947 | 920 | 741 | 712 | 692 |
| 1628 |  | May | 31 | 964 | 923 | 903 | 725 | 694 | 679 |
| 1629 |  | June | 27 | 1055 | 1001 | 948 | 793 | 753 | 713 |
| 1630 |  | July | 24 | 1065 | 996 | 947 | 801 | 749 | 712 |
| 1631 |  | August | 20 | 1109 | 1045 | 1013 | 834 | 786 | 762 |
| 1632 |  | September | 16 | 1003 | 947 | 920 | 754 | 712 | 692 |
| 1633 |  | October | 13 | 1003 | 959 | 935 | 754 | 721 | 703 |
| 1634 |  | November | 9 | 1032 | 971 | 932 | 776 | 730 | 701 |
| 1635 |  | December | 7 | 1045 | 996 | 940 | 786 | 749 | 707 |
| 1636 | 1953 | January | 2 | 1015 | 971 | 922 | 763 | 730 | 693 |
| 1637 |  | January | 29 | 966 | 910 | 895 | 726 | 684 | 673 |
| 1638 |  | February | 25 | 879 | 867 | 852 | 661 | 652 | 641 |
| 1639 |  | March | 24 | 914 | 879 | 859 | 687 | 661 | 646 |
| 1640 |  | April | 21 | 996 | 923 | 879 | 749 | 694 | 661 |
| 1641 |  | May | 17 | 904 | 860 | 846 | 680 | 647 | 636 |
| 1642 |  | June | 13 | 887 | 872 | 863 | 667 | 656 | 649 |
| 1643 |  | July | 10 | 875 | 855 | 843 | 658 | 643 | 634 |
| 1644 |  | August | 6 | 946 | 892 | 863 | 711 | 671 | 649 |
| 1645 |  | September | 2 | 899 | 872 | 858 | 676 | 656 | 645 |
| 1646 |  | September | 29 | 916 | 899 | 887 | 689 | 676 | 667 |
| 1647 | - | October | 26 | 903 | 879 | 874 | 679 | 661 | 657 |
| 1648 |  | November | 22 | 872 | 867 | 855 | 656 | 652 | 643 |
| 1649 |  | December | 19 | 872 | 867 | 862 | 656 | 652 | 648 |
| 1650 | 1954 | January | 15 | 864 | 862 | 850 | 650 | 648 | 639 |
| 1651 |  | February | 11 | 860 | 855 | 855 | 647 | 643 | 643 |

TABLE III. - Day-time and nightotime temperatures at the thermopause. (continued)

| Rota= tion | $\begin{gathered} \text { Date } \\ ( \pm 13 \text { days }) \end{gathered}$ |  |  | Day-time |  |  | Night-time |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  |  | Max. | Average | Min. | Max. | Average | Min. |
| 1652 | 1954 | March | 10 | 918 | 879 | 867 | 690 | 661 | 652 |
| 1653 |  | April | 6 | 860 | 855 | 850 | 647 | 643 | 639 |
| 1654 |  | May | 3 | 851 | 848 | 846 | 640 | 638 | 636 |
| 1655 |  | May | 30 | 848 | 843 | 838 | 638 | 634 | 630 |
| 1656 |  | June | 26 | 839 | 836 | 834 | 631 | 629 | 627 |
| 1657 |  | July | 23 | 858 | 848 | 846 | 645 | 638 | 636 |
| 1658 |  | Aucust | 19 | 874 | 862 | 852 | 657 | 648 | 641 |
| 1659 |  | September | 15 | 867 | 862 | 859 | 652 | 648 | 646 |
| 1660 |  | Octcber | 12 | 902 | 892 | 883 | 678 | 671 | 664 |
| 1661 |  | November | 8 | 910 | 886 | 871 | 684 | 666 | 655 |
| 1662 |  | December | 5 | 910 | 886 | 874 | 684 | 666 | 657 |
| 1663 | 1955 | January | 1 | 1003 | 947 | 912 | 754 | 712 | 686 |
| 1664 |  | January | 28 | 964 | 935 | 908 | 725 | 703 | 683 |
| 1665 |  | February | 24 | 922 | 910 | 900 | 693 | 684 | 677 |
| 1666 |  | March | 23 | 888 | 886 | 883 | 668 | 666 | 664 |
| 1667 |  | April | 19 | 907 | 892 | 883 | 682 | 671 | 664 |
| 1668 |  | May | 16 | 963 | 928 | 904 | 724 | 698 | 680 |
| 1669 |  | June | 12 | 1057 | 984 | 946 | 795 | 740 | 711 |
| 1670 |  | July | 9 | 986 | 959 | 932 | 741 | 721 | 701 |
| 1671 |  | August | 5 | 1067 | 984 | 926 | 802 | 740 | 696 |
| 1672 |  | Seprember | 1 | 1043 | 996 | 940 | 784 | 749 | 707 |
| 1673 |  | September | 28 | 1071 | 1015 | 976 | 805 | 763 | 734 |
| 1674 |  | October | 25 | 1200 | 1106 | 1031 | 902 | 832 | 775 |
| 1675 |  | November | 21 | 1228 | 1198 | 1174 | 923 | 901 | 883 |
| 1676 |  | December | 18 | 1253 | 1198 | 1158 | 942 | 901 | 861 |
| 1677 | 1956 | January | 14 | 1415 | 1283 | 1190 | 1064 | 965 | 895 |
| 1678 |  | February | 10 | 1734 | 1418 | 1237 | 1304 | 1066 | 930 |
| 1679 |  | March | 8 | 1487 | 1382 | 1329 | 1118 | 1039 | 999 |
| 1680 |  | April | 4 | 1470 | 1387 | 1297 | 1105 | 1043 | 975 |

TABLE III.- Day-time and night-time temperatures at the thermopause. (continued)

| Rotation | $\begin{gathered} \text { Date } \\ ( \pm 13 \text { days }) \end{gathered}$ |  |  | Day-time |  | Night-time |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  |  | Max. | Average | Min. | Max. | Average | Min. |
| 1681 | 1956 | May | 1 | 1486 | 1412 | 1305 | 1117 | 1062 | 981 |
| 1682 |  | May | 28 | 1455 | 1382 | 1309 | 1094 | 1039 | 984 |
| 1683 |  | June | 24 | 1462 | 1368 | 1298 | 1099 | 1029 | 976 |
| 1684 |  | July | 21 | 1483 | 1424 | 1351 | 1115 | 1071 | 1016 |
| 1685 |  | August | 17 | 1722 | 1632 | 1541 | 1295 | 1227 | 1159 |
| 1686 |  | September | 13 | 1807 | 1644 | 1553 | 1359 | 1236 | 1168 |
| 1687 |  | October | 10 | 1673 | 1571 | 1483 | 1258 | 1181 | 1115 |
| 1688 |  | November | 6 | 2071 | 1871 | 1705 | 1557 | 1407 | 1282 |
| 1689 |  | December | 3 | 1885 | 1785 | 1670 | . 1417 | 1342 | 1256 |
| 1690 |  | December | 30 | 2068 | 1907 | 1766 | 1555 | 1434 | 1328 |
| 1691 | 1957 | January | 26 | 1782 | 1626 | 1541 | 1340 | 1223 | 1159 |
| 1692 |  | February | 22 | 1523 | 1491 | 1452 | 1145 | 1121 | 1092 |
| 1693 |  | March | 21 | 1595 | 1571 | 15.44 | 1199 | 1181 | 1161 |
| 1694 |  | April | 17 | 1641 | 1565 | 1487 | 1234 | 1177 | 1118 |
| 1695 |  | May | 14 | 1721 | 1613 | 1511 | 1294 | 1213 | 1136 |
| 1696 |  | June | 10 | 1975 | 1785 | 1626 | 1485 | 1342 | 1223 |
| 1697 |  | July | 7 | 2040 | 1780 | 1607 | 1534 | 1338 | 1208 |
| 1698 |  | August | 3 | 1760 | 1596 | 1506 | 1323 | 1200 | 1132 |
| 1699 |  | August | 30 | 1971 | 1810 | 1649 | 1482 | 1361 | 1240 |
| 1700 |  | September | 26 | 2212 | 2048 | 1931 | 1663 | 1540 | 1452 |
| 1701 |  | October | 23 | 2269 | 2140 | 1996 | 1706 | 1609 | 1501 |
| 1702 |  | November | 19 | 1938 | 1841 | 1770 | 1457 | 1384 | 1331 |
| 1703 |  | December | 16 | 2391 | 2072 | 1835 | 1798 | 1558 | 1380 |
| 1704 | 1958 | January | 12 | 2027 | 1871 | 1701 | 1524 | 1407 | 1279 |
| 1705 |  | February | 8 | 1772 | 1620 | 1515 | 1332 | 1218 | 1139 |
| 1706 |  | March | 7 | 1818 | 1700 | 1563 | 1368 | 1278 | 1175 |
| 1707 |  | April | 3 | 2264 | 2006 | 1730 | 1702 | 1508 | 1301 |
| 1708 |  | April | 30 | 1904 | 1797 | 1685 | 1432 | 1351 | 1267 |


| Rotation | $\begin{gathered} \text { Date } \\ ( \pm 13 \text { days }) \end{gathered}$ |  |  | Day-time |  |  | Night-time |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  |  | Max. | Average | Min. | Max. | Average | Min. |
| 1709 | 1958 | May | 27 | 1749 | 1620 | 1549 | 1315 | 1218 | 1165 |
| 1710 |  | June | 23 | 1721 | 1638 | 1541 | 1294 | 1232 | 1159 |
| 1711 |  | July | 20 | 1887 | 1700 | 1583 | 1419 | 1278 | 1190 |
| 1712 |  | August | 16 | 1835 | 1736 | 1680 | 1380 | 1305 | 1263 |
| 1713 |  | September |  | 1948 | 1834 | 1744 | 1465 | 1379 | 1311 |
| 1714 |  | October | 9 | 1920 | 1730 | 1596 | 1444 | 1301 | 1200 |
| 1715 |  | November | 5 | 1709 | 1577 | 1472 | 1285 | 1186 | 1107 |
| 1716 |  | December | 2 | 1938 | 1853 | 1660 | 1457 | 1393 | 1248 |
| 1717 |  | December | 29 | 1983 | 1827 | 1713 | 1491 | 1374 | 1288 |
| 1718 | 1959 | January | 25 | 2246 | 1944 | 1734 | 1689 | 1462 | 1304 |
| 1719 |  | February | 21 | 1698 | 1613 | 1531 | 1277 | 1213 | 1151 |
| 1720 |  | March | 20 | 1998 | 1834 | 1661 | 1502 | 1379 | 1249 |
| 1721 |  | April | 16 | 1636 | 1577 | 1519 | 1230 | 1186 | 1142 |
| 1722 |  | May | 13 | 1874 | 1669 | 1576 | 1409 | 1255 | 1185 |
| 1723 |  | June | 9 | 1722 | 1632 | 1530 | 1295 | 1227 | 1150 |
| . 1724 |  | July | 6 | 1869 | 1669 | 1532 | 1405 | 1255 | 1152 |
| 1725 |  | August | 2 | 1692 | 1577 | 1495 | 1272 | 1186 | 1124 |
| 1726 |  | August | 29 | 2077 | 1846 | 1700 | 1562 | 1388 | 1278 |
| 1727 |  | September | 25 | 1480 | 1387 | 1326 | 1113 | 1043 | 997 |
| 1728 |  | October | 22 | 1452 | 1394 | 1350 | 1092 | 1048 | 1015 |
| 1729 |  | November | 18 | 1685 | 1522 | 1410 | 1267 | 1144 | 1060 |
| 1730 |  | December | 15 | 1623 | 1491 | 1423 | 1220 | 1121 | 1070 |
| 1731 | 1960 | January | 11 | 1636 | 1516 | 1434 | 1230 | 1140 | 1078 |
| 1732 |  | February | 7 | 1708 | 1552 | 1415 | 1284 | 1167 | 1064 |
| 1733 |  | March | 5 | 1252 | 1234 | 1208 | 941 | 928 | 908 |
| 1734 |  | April | 1 | 1539 | 1424 | 1325 | 1157 | 1071 | 996 |
| 1735 |  | April | 28 | 1418 | 1357 | 1301 | 1066 | 1020 | 978 |
| 1736 |  | May | 25 | 1443 | 1382 | 1350 | 1085 | 1039 | 1015 |

TABLE III.- Day-time and night-time temperatures at the thermopause. (continued)

| Rota- <br> tion | $\begin{gathered} \text { Date } \\ ( \pm 13 \text { days }) \end{gathered}$ |  |  | Day-time |  |  | Night-time |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  |  | Max. | Average | Min. | Max. | Average | Min. |
| 1737 | 1960 | June | 21 | 1531 | 1375 | 1275 | 1151 | 1034 | 959 |
| 1738 |  | July | 18 | 1526 | 1345 | 1283. | 1147 | 1011 | 965 |
| 1739 |  | August | 14 | 1664 | 1486 | 1322 | 1251 | 1117 | 994 |
| 1740 |  | September | 11 | 1472 | 1382 | 1282 | 1107 | 1039 | 964 |
| 1741 |  | October | 7 | 1338 | 1277 | 1194 | 1006 | 960 | 898 |
| 1742 |  | November | 3 | 1514 | 1338 | 1245 | 1138 | 1006 | 936 |
| 1743 |  | November | 30 | 1339 | 1246 | 1168 | 1007 | 937 | 878 |
| 1744 |  | December | 27 | 1357 | 1222 | 1137 | 1020 | 919 | 855 |
| 1745 | 1961 | January | 23 | 1141 | 1088 | 1049 | 858 | 818 | 789 |
| 1746 |  | February | 19 | 1073 | 1039 | 1021 | 807 | 781 | 768 |
| 1747 |  | March | 18 | 1112 | 1051 | 1019 | 836 | 790 | 766 |
| 1748 |  | April | 14 | 1106 | 1045 | 1011 | 832 | 786 | 760 |
| 1749 |  | May | 11 | 1097 | 1039 | 1000 | 825 | 781 | 752 |
| 1750 |  | June | 7 | 1150 | 1057 | 1008 | 865 | 795 | 758 |
| 1751 |  | July | 4 | 1204 | 1130 | 1072 | 905 | 850 | 806 |
| 1752 |  | July | 31 | 1140 | 1081 | 1020 | 857 | 813 | 767 |
| 1753 |  | August | 27 | 1125 | 1093 | 1052 | 846 | 822 | 791 |
| 1754 |  | September | 23. | 1178 | 1093 | 1047 | 886 | 822 | 787 |
| 1755 |  | October | 20 | 1056 | 1015 | 983 | 794 | 763 | 739 |
| 1756 |  | November | 16 | 1025 | 984 | 958 | 771 | 740 | 720 |
| 1757 |  | December | 13 | 1055 | 1008 | 967 | 793 | 758 | 727 |
| 1758 | 1962 | January | 9 | 1048 | 978 | 9.46 | 788 | 735 | 711 |
| 1759 |  | February | 5 | 1094 | 1039 | 990 | 823 | 781 | 744 |
| 1760 |  | March | 4 | 1177 | 1063 | 990 | 885 | 799 | 744 |
| 1761 |  | March | 31 | 1129 | 1039 | 978 | 849 | 781 | 735 |
| 1762 |  | April | 27 | 1085 | 1039 | 990. | 816 | 781 | 744 |
| 1763 |  | May | 24 | 1076 | 1032 | 991 | 809 | 776 | 745 |
| 1764 |  | June | 20 | 998 | 978 | 966 | 750 | 735 | 726 |


| Rotation | $\begin{gathered} \quad \text { Date } \\ ( \pm 13 \text { days }) \end{gathered}$ |  |  | Day-time |  |  | Night-time |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  |  | Max. | Average | Min. | Max. | Average | Min. |
| 1765 | 1962 | July | 17 | 943 | 923 | 899 | 709 | 694 | 676 |
| 1766 |  | August | 13 | 939 | 904 | 887 | 706 | 680 | - 66.7 |
| 1767 |  | September | 9 | 1015 | 971 | 924 | 763 | 730 | 695 |
| 1768 |  | October | 6 | 1000 | 971 | 954 | 752 | 730 | 717 |
| 1769 |  | November | 2 | 986 | 959 | 942 | 741 | 721 | 708 |
| 1770 |  | November | 29 | 984 | 940 | 926 | 740 | 707 | 696 |



Fig. 16.- Average night-time temperatures at the thermopause with minimum-to-maximum values during ——a 27-day period.


Fig. 17.- Average daytime temperatures at the thermopause with minimum-tomaximum values during a — 27-day period.

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