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# Variability of Martian water ice and dust vertical profiles with TGO/NOMAD

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#### Abstract

We present the first results of water ice and dust microphysical properties derived from the Nadir and Occultation for Mars Discovery (NOMAD) instrument, onboard ExoMars Trace Gas Orbiter (TGO). These retrievals are derived from vertical profiles that span the entire science phase so far, which began April 21st, 2018.

#### 1. NOMAD instrument

The NOMAD spectrometer onboard ExoMars TGO operates in the wavelength range between 0.2-4.3 µm [1]. This instrument suite has three spectral channels: a solar occultation channel (SO – Solar Occultation; 2.3-4.3 µm), a second infrared channel capable of nadir, solar occultation, and limb sounding (LNO -Limb Nadir and solar Occultation; 2.3-3.8 µm), and an ultraviolet/visible channel (UVIS - UV visible, 200-650 nm). The infrared channels (SO and LNO) have high resolving power ( $\lambda/d\lambda \sim 10,000-20,000$ ) provided by an echelle grating in combination with an Acousto Optic Tunable Filter (AOTF), which restricts the bandpass width and selects the observed diffraction order [2]. The infrared channels were derived from the Solar Occultation in the IR (SOIR) instrument [3] onboard Venus Express. The high sampling rate (1 second) of SO measurements provides unprecedented vertical resolution (~ 1 km), which span altitudes from the surface to ~200 km. Moreover, the instantaneous change of the observed diffraction orders, made possible by the AOTF, enables the SO channel to measure five or six different diffraction orders per second.

In this study, we combine broadband (2700 to 4300 cm<sup>-1</sup>) information in single occultations to infer the

vertical structure of dust and water ice profiles. The sensitivity and accuracy of this kind of analysis is maximized when the diffraction orders observed are not heavily affected by molecular extinction due to CO<sub>2</sub>. The characterization of the sensitivity of each diffraction order to dust and water ice is performed as a preliminary step to the retrievals, according to the calibration scheme developed for the NOMAD instrument [4]. The combined broadband transmittance (I/F) has been used to infer the dust and water ice concentrations and their particle size, for each altitude from 0 to 100 km in each occultation. Data have been analyzed with the Planetary Spectrum Generator (PSG [6]) developed at NASA GSFC, which embodies a full retrieval package based on a state-of-art implementation of Optimal Estimation [7].

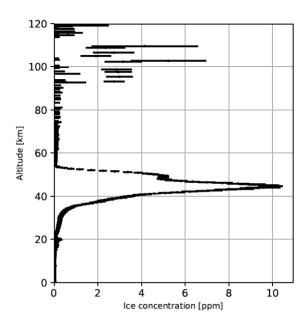
# 2. Water ice vertical profiles

Using the data acquired by NOMAD from April 2018 to April 2019, we have analyzed the vertical distribution of water ice from the surface up to 110 km, as well as the vertical variation of particle size. The measurements acquired between June and September 2018, coincide with a Planet Encircling Dust Event (PEDE). The effect of this storm on the properties and localization of water ice clouds is presented and discussed in depth. We find that high altitude (mesospheric) water ice clouds are detected frequently and that their microphysical characteristics differ from low-altitude clouds. Their altitude has also been observed to abruptly increase as the PEDE starts, and the velocity at which this happens (of the order of a Martian day) can provide feedback about the dynamics behind high altitude cloud formation. We have also characterized the latitudinal distributions of water ice clouds, and we elaborate on the implications of these observations to the advance of the current knowledge of the water cycle on Mars [5].

## 3. Dust vertical profiles

We have simultaneously retrieved dust vertical profiles of concentration and average particle sizes for this same dataset. We discuss the temporal evolution and morphology of dust profiles, with particular attention to the PEDE and a second, less intense dust storm in January-February 2019. This provides evidence for the thermal effects of dust on the atmosphere, and its role in cloud formation processes. In particular, we have found that dust can lift up in the atmosphere to altitudes ~60 km during dust storms, forming suspended layers with complex vertical structures.

## **Figure**



**Figure 1**: Example of water ice vertical profile retrieved from the NOMAD-SO data before the global dust storm in 2018 ( $L_S = 171.45^{\circ}$  and latitude  $43^{\circ}$  N– $68^{\circ}$  N) [5].

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