



No evidence of phosphine in the atmosphere of Venus from independent analyses

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The detection of phosphine (PH₃) in the atmosphere of Venus has been recently reported on the basis of millimetre-wave radio observations¹ and their reanalyses^{2,3}. In this Matters Arising we perform an independent reanalysis, identifying several issues in the interpretation of the spectroscopic data. As a result, we determine sensitive upper limits for PH₃ in Venus's atmosphere (>75 km, above the cloud decks) that are discrepant with the findings in refs. 1–3.

The measurements target the fundamental first rotational transition of PH₃ ($J=1-0$) at 266.944513 GHz, which was observed with the James Clerk Maxwell Telescope (JCMT) in June 2017 and with the Atacama Large Millimeter/submillimeter Array (ALMA) in March 2019. This line's centre is near the SO₂ ($J=30_{9,21}-31_{8,24}$) transition at 266.943329 GHz (only 1.3 km s⁻¹ away from the PH₃ line), which represents a potential source of contamination. The JCMT and ALMA data, as presented in ref. 1, are at spectral resolutions comparable to the frequency separation of the two lines. Moreover, the spectral features identified are several kilometres per second in width, and therefore do not permit distinct spectroscopic separation of the candidate spectral lines of PH₃ and SO₂. We present the radiative transfer modelling we have performed and then discuss the ALMA and JCMT analyses in turn.

ALMA reanalysis

The analysis of interferometric data is relatively complex, in particular for bright and extended sources such as Venus (15.2 arc-sec angular diameter for the ALMA data). The completeness of the different baselines (short and long) determines the ability to accurately measure the total planetary flux density⁴, while the bandpass calibration is a crucial factor in the ultimate quality of the resulting spectra^{5,6}. The details of the calibration and imaging used during data reduction can have a dramatic impact on the quality and validity of the resulting ALMA interferometric data. The extracted spectra in extended data fig. 4 and the interferometric map in extended data fig. 3 of ref. 1 show large quasiperiodic

fluctuations in the spectrum, which in ref. 1 is fitted with high-order polynomials. Particularly challenging is the fact that these fluctuations have a pattern/width comparable to their defined PH₃ line core region. As we present in Supplementary Section 2 and as also reported in refs. 7,8, artificially produced features that mimic true atmospheric lines can be produced when analysing data with such characteristics.

Many of these large fluctuations can be introduced by the particular parameters used in calibration during data reduction. After the publication of ref. 1 we notified the North American ALMA Science Center team of substantial differences in the final spectra when applying different treatments of the bandpass calibrations—specifically, the type of bandpass calibration solution (traditional channel to channel versus polynomial fitted versus smoothed), and in particular enabling the ‘usescratch=True’ setting in Common Astronomy Software Applications (CASA)'s setjy for the model used for Callisto (the bandpass calibrator). The newly reprocessed data from the Joint ALMA Office (JAO) take these issues into consideration and are used by refs. 2,3. We analysed the data as presented originally in ref. 1 and could not recover a PH₃ signature (Fig. 1b), and we also reduced the ALMA data using the new JAO scripts as the starting point. We employed three separate procedures by three different groups (National Aeronautics and Space Administration Goddard Space Flight Center—NASA/GSFC, Berkeley, and National Radio Astronomy Observatory—NRAO), making use of different methods to evaluate the significance of the reported ALMA PH₃ signatures. The three methods are the following: (1) employing the updated JAO scripts throughout (using CASA only, done at NASA/GSFC); (2) employing the updated JAO scripts, but proceeding independently beyond self-calibration (using CASA only; done at Berkeley)—note that phase-only self-calibration was done but not amplitude self-calibration, because the latter can be problematic for extended sources⁷; (3) as (2), but with the post-JAO-script reduction done in the Astronomical Image Processing System (AIPS)

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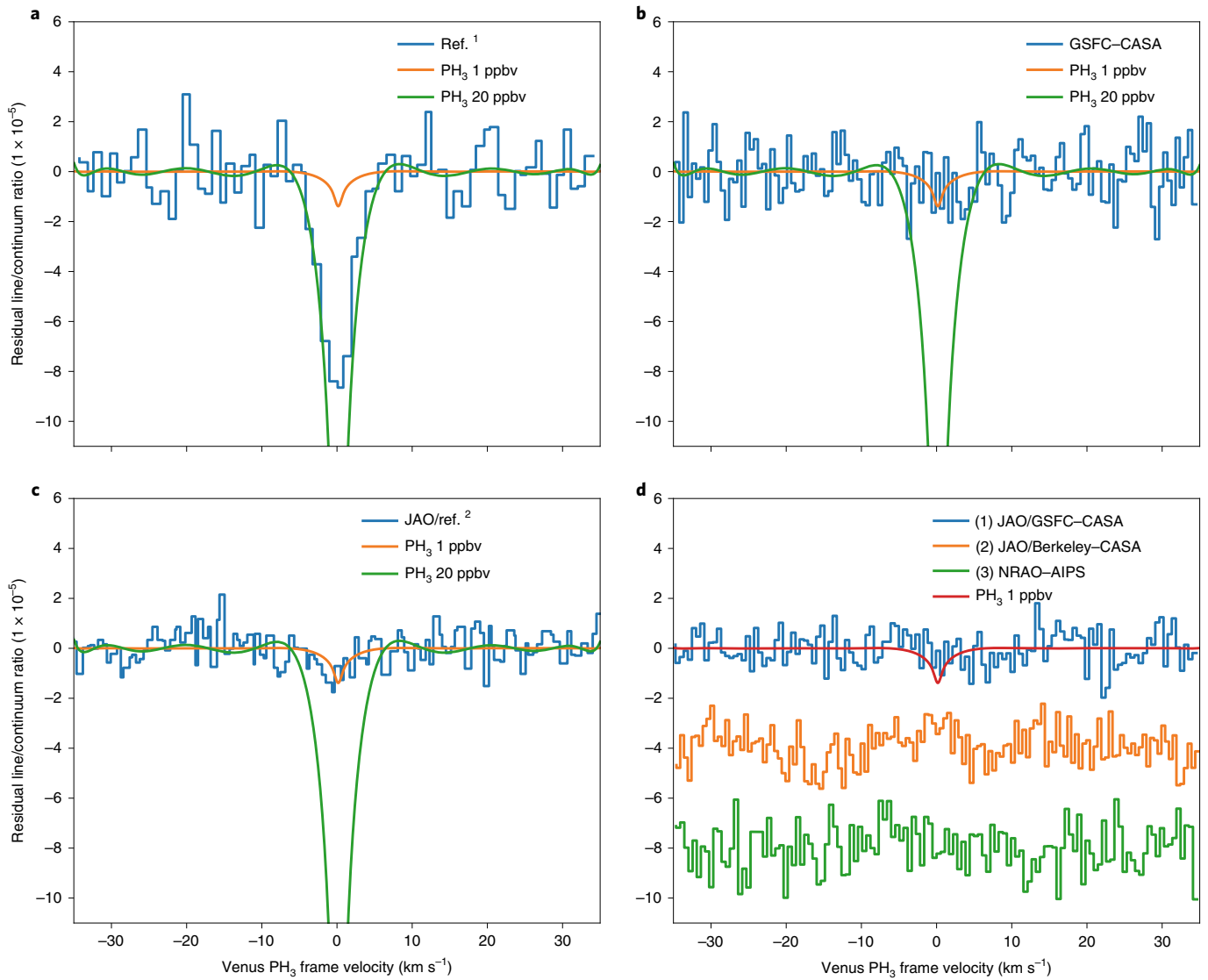


Fig. 1 | Comparison between the results obtained by the authors of the original paper and its reanalysis and our independent analysis of the same data retrieved from the ALMA science archive. a, The ALMA data as presented in fig. 2 of ref. ¹ for the whole planet (the spectral resolution Δv is 1 km s^{-1}). **b**, Our original analysis of the raw ALMA data (Δv : 0.55 km s^{-1}), employing the original ref. ¹ scripts but enabling the `uscratch=True` setting in CASA's `setJy`, correcting the bandpass calibration. **c**, The reanalysed data presented in ref. ² using updated scripts; ref. ¹ employed a higher (12th-order) polynomial in the bandpass calibration (instead of third order in the JAO script), and excluded short baselines ($<33 \text{ m}$) from their analysis (Δv : 0.55 km s^{-1}). The orange and green lines in **a–c** are models of the phosphine line computed with a constant abundance of 1 and 200 ppbv respectively. **d**, Our independent analysis of the JAO-revised ALMA data, employing different methods, a resolution of Δv : 0.55 km s^{-1} , methods 1 and 2 including all baselines and method 3 excluding $<33 \text{ m}$ baselines. A sixth-order polynomial was removed for (1) (see details in Methods and in Supplementary Section 2), while only a second was removed for (2) and (3). The red line corresponds to the model with a constant abundance of 1 ppbv PH_3 .

(using both CASA and AIPS, done at NRAO). We investigated including/removing certain baselines; further details of the methods can be found in Methods and in Supplementary Section 2. All these analyses used the common foundation of the revised JAO scripts to carry out the initial calibration (in particular the bandpass and complex gain versus time calibrations), and very similar steps in the further data reduction (phase-only self-calibration, continuum subtraction, forming the image cube and extracting the disk-averaged spectrum).

The quality of the bandpass calibration was improved with the updated JAO scripts, yet the residual spectrum still showed notable large fluctuations (we found that a sixth-order polynomial captured most of the residual large fluctuations for method 1; ref. ¹ considered a 12th order). Our other two reanalyses (2 and 3) led to practi-

cally flat spectra only needing second-order polynomial baselines (more information on the origin of these differences is explained in Supplementary Section 2). Ultimately, all our analyses of the data using these different approaches and methodologies reveal no conclusive signature of PH_3 (Fig. 1), leading to an upper limit of <1 part per billion by volume (ppbv) (3σ when employing a linewidth of $0.186 \text{ cm}^{-1} \text{ atm}^{-1}$). When employing other proposed linewidths, the upper limit would be $\text{PH}_3 < 0.7 \text{ ppbv}$ (linewidth of $0.12 \text{ cm}^{-1} \text{ atm}^{-1}$) or $<1.5 \text{ ppbv}$ (linewidth of $0.286 \text{ cm}^{-1} \text{ atm}^{-1}$). This upper limit is also consistent with other recently reported upper limits derived from infrared ground-based observations ($\text{PH}_3 < 5 \text{ ppbv}$)⁹ and from spacecraft data ($\text{PH}_3 < 0.2 \text{ ppbv}$)¹⁰. We further validated our analysis of the ALMA data by analysing other SO_2 and HDO lines in the same dataset (Supplementary Section 3).

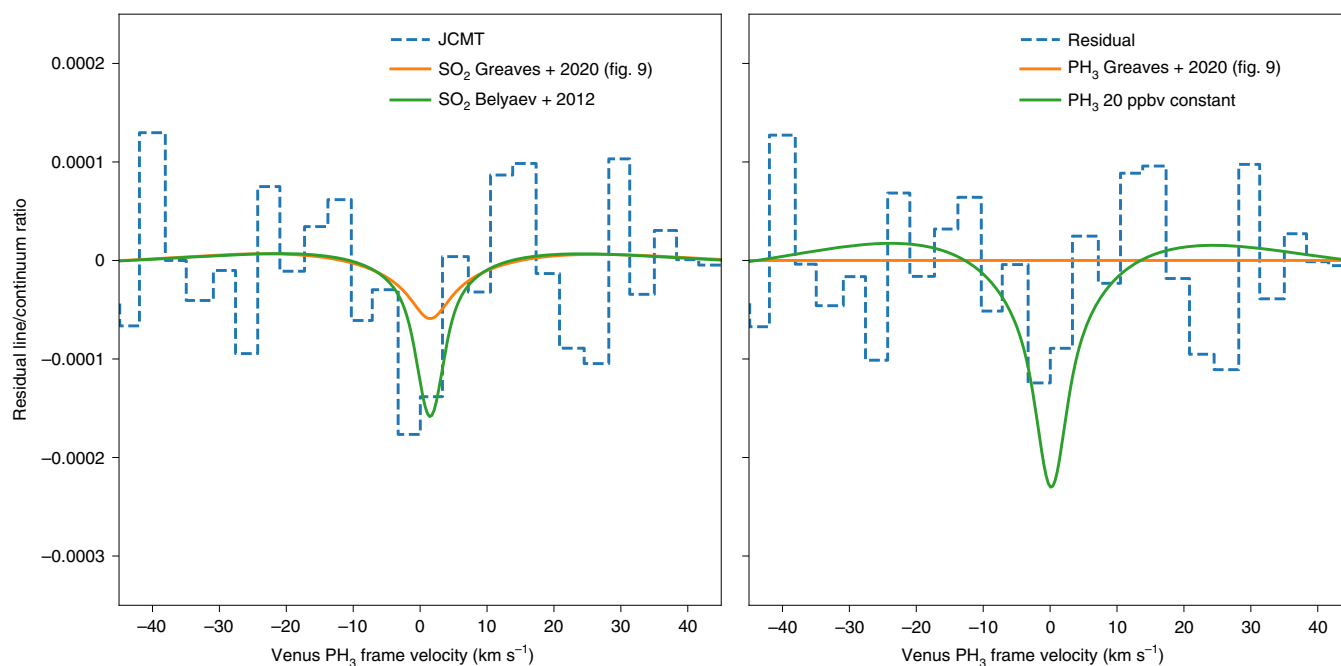


Fig. 2 | Residual JCMT data and models of SO₂ and PH₃. Left: JCMT data for their mid-range solution with masking within ± 5 km s⁻¹, while ‘SO₂ Greaves + 2020’ is a model spectrum synthesized using the T - P in ref. ¹ fig. 8 and the SO₂ profile in ref. ¹ fig. 9. ‘SO₂ Belyaev + 2012’ is a typical profile in ref. ¹¹ (fig. 9) and similar to ref. ¹² (case D), with an SO₂ abundance of ~ 30 ppbv at 80 km, increasing to ~ 100 ppbv at 90 km and reaching ~ 300 ppbv at 95 km. Right: ‘Residual’ is obtained by removing the SO₂ signature from the JCMT data as modelled with a 3 km s⁻¹ resolution using the ref. ¹ fig. 9 SO₂ profile. The ‘PH₃ Greaves + 2020’ spectrum is a model spectrum employing the PH₃ profile in ref. ¹ fig. 9 (peaking at 60 km with 20 ppbv) and the default linewidth (0.186 cm⁻¹ atm⁻¹), while a model considering a vertically constant mixing ratio of 20 ppbv is labelled ‘PH₃ 20 ppbv constant’. ‘PH₃ Greaves + 2020’ is flat and featureless because these measurements only sample PH₃ above 75 km (Supplementary Section 4).

JCMT analysis and SO₂ contamination

The analysis of JCMT spectral data in ref. ¹ involved fitting of multiple polynomials for the purpose of continuum subtraction: an initial fourth-order polynomial was removed, then a ninth-order polynomial was removed from a smoothed spectrum and finally a masked eighth-order polynomial was removed. Recently, Thompson⁸ explored the robustness of the two detection methods used in ref. ¹, namely low-order polynomial fits and higher-order multiple polynomial fits, and found that neither line detection method is able to recover a statistically significant detection at the position of the PH₃/SO₂ line.

We investigate whether a weak PH₃ signature, if present, could be distinguished from SO₂ contamination, and what its contribution to the detected signal could be. To explore this, we employed the same VIR45 temperature–pressure (T - P) profile as used by Greaves et al.¹ (their extended data fig. 8) and the SO₂ profile presented in their extended data fig. 9 (~ 100 ppbv in the 70–90 km region, dropping to < 30 ppbv above 95 km). If this potential SO₂ contaminant signature were removed from the JCMT data, then the original feature would be confined within the noise (Fig. 2). We next modelled a mesospheric SO₂ profile as measured by Venus Express using solar occultation¹¹, and similar to case D of ref. ¹² (SO₂ abundance of ~ 30 ppbv at 80 km, increasing to ~ 100 ppbv at 90 km and reaching ~ 300 ppbv at 95 km). In this case, the SO₂ contamination signature is even stronger and fully captures the claimed PH₃ residual in refs. ^{1–3} (Fig. 2). A comprehensive compilation of SO₂ measurements^{7,8}, including thousands of measurements from the Venus Express orbiter at ultraviolet and infrared wavelengths as well as ground-based observations^{11,13–22}, show that SO₂ abundances at 80 km are frequently well in excess of 100 ppbv, with peak abundances occasionally exceeding 1,000 ppbv. This shows that the SO₂ profiles we adopted are well within usual ranges of variability observed on Venus (more information about the considered profiles can be found in Supplementary Section 1). Having shown that the

absorption line can be reproduced by a reasonable vertical profile of SO₂, we therefore argue that this absorption line cannot be definitively attributed to PH₃.

Probing altitude

We also explored the altitudes from which these absorptions originate, which can be used to constrain photochemical models and facilitate comparison with other measurements^{23–27}. This is ultimately related to the spectroscopic parameters of the targeted line and of the competing radiative active species in this spectral region. When considering the linewidth of 0.186 cm⁻¹ atm⁻¹ for PH₃, at 70 km (3.4×10^{-2} atm) the line would be 213 km s⁻¹ wide, that is, much broader than the narrow window region of ± 5 km s⁻¹ or ± 10 km s⁻¹ searched for PH₃. As a polynomial is fitted and subtracted, removing any spectral line information beyond this core region, the spectral information at broader widths (originating from lower altitudes) is thus removed. Our models predict an observable PH₃ absorption at this frequency only when PH₃ is present above 75 km (see details in Supplementary Section 4), and therefore these data provide no constraints on its abundance in the cloud deck (50–70 km), in contrast to refs. ^{1–3}. This can be seen in Fig. 2: a flat spectrum is calculated for the PH₃ profile determined from the chemical modelling described in ref. ¹, in which PH₃ is present only below 70 km altitude.

Conclusions

By our independent analysis of the ALMA data, we set a limit of 1 ppbv for PH₃. We also show that the observed JCMT feature could be attributed to mesospheric SO₂ gas. Furthermore, for any PH₃ signature to be recoverable in either ALMA or JCMT analyses, PH₃ needs to be present at altitudes above 75 km. We conclude that the recent identification of PH₃ in the upper atmosphere of Venus is not supported.

Methods

We have employed three independent radiative transfer models: the Planetary Spectrum Generator (PSG; <https://psg.gsfc.nasa.gov>)²⁸, the Non-linear Optimal Estimator for Multivariate Spectral Analysis (NEMESIS; <https://nemesiscode.github.io>)²⁹ and the Center for Astrophysics planetary modelling tool³⁰. The Planetary Spectrum Generator radiative transfer analysis included the latest HITRAN SO₂ line parameters for a CO₂ atmosphere³¹, a layer-by-layer, line-by-line study, and a full disk sampling scheme with 10 concentric rings. The NEMESIS analysis was also performed in line-by-line mode, and used the same spectroscopic data with a five-point Gauss–Lobatto disk integration scheme. As described in ref. ¹, there is some uncertainty in the line-shape parameters for the PH₃ line in a CO₂ atmosphere. HITRAN reports an air linewidth of 0.067 cm⁻¹ atm⁻¹, which would correspond to 0.12 cm⁻¹ atm⁻¹ in a CO₂ atmosphere if the value were scaled by the typical 1.8 scaling ratio observed for the SO₂ lines³². Greaves et al.¹ considered a theoretical estimate of 0.186 cm⁻¹ atm⁻¹, and an upper range of 0.286 cm⁻¹ atm⁻¹ measured for the NH₃(*J* = 1–0) line in CO₂ at 572.498160 GHz. Considering the uncertainty on this parameter, we adopt the same value as Greaves et al.¹ by default (0.186 cm⁻¹ atm⁻¹).

When reducing the ALMA data, we downloaded the raw data and scripts as provided in the archive. By three separate reduction methods, we then used the provided scripts along with well established techniques in the calibration and imaging of radio interferometric data to obtain final disk-averaged spectra for Venus. There are three scripts provided with the data: (1) a calibration script; (2) an imaging preparation script; (3) an imaging script. The calibration script calculates and applies several system calibrations and flags data known to be invalid, then sets the model for Callisto and uses this for flux density scale and bandpass calibration (using a third-order polynomial fit to the bandpass amplitude), then carries out a complex gain versus time calibration using the calibrator J2000–1748. The imaging preparation script Doppler shifts the spectral axis of the various spectral windows, then images Venus and uses that image to carry out phase-only self-calibration, then sets the model for Venus and carries out amplitude self-calibration.

The imaging script then makes a continuum image, carries out continuum subtraction and makes the image cube. In both the continuum image and the image cube a primary beam correction is made, since the primary beam at these frequencies (with a full-width at half-maximum of ~24") resolves the disk of Venus (diameter ~15.2" at the time). We note that the amplitude self-calibration step in the imaging preparation script can lead to problems when observing bright objects with sharp boundaries (such as Venus or Jupiter)³, and it also leads to the flux density scale being set incorrectly when using these scripts (because it will set the model flux density to that expected for Venus with no primary beam attenuation, but the primary beam actually attenuates that flux density by about 15%, and this attenuation is actually corrected for in the imaging script). All three of our independent analyses used the entire calibration script as provided from the archive. Our first analysis continued with the scripts as provided from the archive throughout, attempting to recreate as closely as possible the analysis in refs. ^{1–3}. Our second analysis took the data after the Doppler shift step, then proceeded with the reduction within CASA. The only notable difference from the imaging preparation and imaging scripts is the use of a limb-darkened model of Venus for the phase-only self-calibration, and not carrying out amplitude self-calibration at all. Our third analysis was nearly identical to the second, taking the data after the Doppler shift step, but then proceeding within AIPS instead of CASA. The differences in this third analysis were addition of a second bandpass calibration using a model of Callisto, use of the same limb-darkened model of Venus as in the second analysis for the phase-only self-calibration, no use of amplitude self-calibration, and a scaling of the flux densities by the ratio of the expected (primary beam-attenuated) flux density from Venus to that calculated with a fit to the visibilities. Aside from these differences, other steps in our second and third analyses were essentially the same as those in the provided scripts. For the third analysis, baselines with separations less than 33 m were excluded when making the final image cube; for the first analysis, this was investigated and found to reduce the amplitude of the bandpass calibration artifacts, while increasing the (channel-to-channel) spectral noise. No PH₃ detection was evident in either case. For all three analyses, we used the entire disk of Venus to create the final spectrum.

Data availability

This paper makes use of the 2018.A.00023.S ALMA data, available at <https://almascience.nrao.edu/asax>. The JCMT data are available at <https://www.eaobservatory.org/jcmt/science/archive>.

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References

- Greaves, J. S. et al. Phosphine gas in the cloud decks of Venus. *Nat. Astron.* <https://doi.org/10.1038/s41550-020-1174-4> (2020).

- Greaves, J. S. et al. Reply to: No evidence of phosphine in the atmosphere of Venus from independent analyses. *Nat. Astron.* <https://doi.org/10.1038/s41550-021-01424-x> (2021).
- Greaves, J. S. et al. Addendum: Phosphine gas in the cloud decks of Venus. *Nat. Astron.* <https://doi.org/10.1038/s41550-021-01423-y> (2021).
- de Pater, I. et al. First ALMA millimeter-wavelength maps of Jupiter, with a multiwavelength study of convection. *Astron. J.* **158**, 139 (2019).
- Thelen, A. E. et al. Detection of CH₃C₃N in Titan's atmosphere. *Astrophys. J. Lett.* **903**, L22 (2020).
- Yamaki, H., Kamenno, S., Beppu, H., Mizuno, I. & Imai, H. Optimization by smoothed bandpass calibration in radio spectroscopy. *Publ. Astron. Soc. Jpn* **64**, 118 (2012).
- Snellen, I. A. G., Guzman-Ramirez, L., Hogerheijde, M. R., Hygate, A. P. S. & van der Tak, F. F. S. Re-analysis of the 267-GHz ALMA observations of Venus: no statistically significant detection of phosphine. *Astron. Astrophys.* **644**, L2 (2020).
- Thompson, M. A. The statistical reliability of 267-GHz JCMT observations of Venus: no significant evidence for phosphine absorption. *Mon. Not. R. Astron. Soc. Lett.* **501**, L18–L22 (2020).
- Encrenaz, T. et al. A stringent upper limit of the PH₃ abundance at the cloud top of Venus. *Astron. Astrophys.* **643**, L5 (2020).
- Trompet, L. et al. Phosphine in Venus' atmosphere: detection attempts and upper limits above the cloud top assessed from the SOIR/VEx spectra. *Astron. Astrophys.* **645**, L4 (2021).
- Belyaev, D. A. et al. Vertical profiling of SO₂ and SO above Venus' clouds by SPICAV/SOIR solar occultations. *Icarus* **217**, 740–751 (2012).
- Lincowski, A. P. et al. Claimed detection of PH₃ in the clouds of Venus is consistent with mesospheric SO₂. Preprint at <https://arxiv.org/abs/2101.09837> (2021).
- Krasnopolsky, V. A. Spatially-resolved high-resolution spectroscopy of Venus 2. Variations of HDO, OCS, and SO₂ at the cloud tops. *Icarus* **209**, 314–322 (2010).
- Sandor, B. J., Todd Clancy, R., Moriarty-Schieven, G. & Mills, F. P. Sulfur chemistry in the Venus mesosphere from SO₂ and SO microwave spectra. *Icarus* **208**, 49–60 (2010).
- Piccialli, A. et al. Mapping the thermal structure and minor species of Venus mesosphere with ALMA submillimeter observations. *Astron. Astrophys.* **606**, A53 (2017).
- Vandaele, A. C. et al. Sulfur dioxide in the Venus atmosphere: I. Vertical distribution and variability. *Icarus* **295**, 16–33 (2017).
- Vandaele, A. C. et al. Sulfur dioxide in the Venus atmosphere: II. Spatial and temporal variability. *Icarus* **295**, 1–15 (2017).
- Vandaele, A. C. Composition and chemistry of the neutral atmosphere of Venus. *Oxford Research Encyclopedia of Planetary Science* <https://oxfordre.com/planetaryscience/view/10.1093/acrefore/9780190647926.001.0001/acrefore-9780190647926-e-4> (2020).
- Encrenaz, T., Moreno, R., Moullet, A., Lellouch, E. & Fouchet, T. Submillimeter mapping of mesospheric minor species on Venus with ALMA. *Planet. Space Sci.* **113–114**, 275–291 (2015).
- Encrenaz, T. et al. HDO and SO₂ thermal mapping on Venus—IV. Statistical analysis of the SO₂ plumes. *Astron. Astrophys.* **623**, A70 (2019).
- Marcq, E. et al. Climatolgy of SO₂ and UV absorber at Venus' cloud top from SPICAV-UV nadir dataset. *Icarus* **335**, 113368 (2020).
- Encrenaz, T. et al. HDO and SO₂ thermal mapping on Venus—V. Evidence for a long-term anti-correlation. *Astron. Astrophys.* **639**, A69 (2020).
- Mogul, R., Limaye, S. S. & Way, M. J. Venus' mass spectra show signs of disequilibria in the middle clouds. Preprint at <https://doi.org/10.1002/essoar.10504552.1> (2020).
- Mogul, R., Limaye, S. S., Way, M. J. & Cordova, J. Is phosphine in the mass spectra from Venus' clouds? Preprint at <https://arxiv.org/abs/2009.12758> (2020).
- Seiff, A. et al. Models of the structure of the atmosphere of Venus from the surface to 100 kilometers altitude. *Adv. Space Res.* **5**, 3–58 (1985).
- Limaye, S. S. et al. Venus atmospheric thermal structure and radiative balance. *Space Sci. Rev.* **214**, 102 (2018).
- Akins, A. B., Lincowski, A. P., Meadows, V. S. & Steffes, P. G. Complications in the ALMA detection of phosphine at Venus. *Astrophys. J.* **907**, L27 (2021).
- Villanueva, G. L., Smith, M. D., Protospapa, S., Faggi, S. & Mandell, A. M. Planetary Spectrum Generator: an accurate online radiative transfer suite for atmospheres, comets, small bodies and exoplanets. *J. Quant. Spectrosc. Radiat. Transf.* **217**, 86–104 (2018).
- Irwin, P. G. J. et al. The NEMESIS planetary atmosphere radiative transfer and retrieval tool. *J. Quant. Spectrosc. Radiat. Transf.* **109**, 1136–1150 (2008).
- Gurwell, M. A., Bergin, E. A., Melnick, G. J. & Tolls, V. Mars surface and atmospheric temperature during the 2001 global dust storm. *Icarus* **175**, 23–31 (2005).
- Gordon, I. E. et al. The HITRAN2016 molecular spectroscopic database. *J. Quant. Spectrosc. Radiat. Transf.* **203**, 3–69 (2017).

32. Wilzewski, J. S., Gordon, I. E., Kochanov, R. V., Hill, C. & Rothman, L. S. H., He, and CO₂ line-broadening coefficients, pressure shifts and temperature-dependence exponents for the HITRAN database. Part 1: SO₂, NH₃, HF, HCl, OCS and C₂H₂. *J. Quant. Spectrosc. Radiat. Transf.* **168**, 193–206 (2016).

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Author contributions

G.L.V., P.G.J.I., M.G., V.K. and G.L. performed the retrievals and the radiative-transfer modelling. M.C., I.d.P. and B.B. calibrated and analysed the ALMA data. S.N.M.,

C.A.N., S.H.L.-C., R.C., A.E.T., A.M., E.M.M., S.C., N.B. and K.R.d.K. assisted with the interpretation of the interferometric spectra. C.F.W., S.F., T.J.F., M.L., P.H., G.N.A., A.M.M., A.C.V. and R.K. assisted with the interpretation of the results in the context of the Venusian atmosphere and its photochemistry. All authors contributed to writing and revising the manuscript.

Competing interests

The authors declare no competing interests.

Additional information

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